

Chapter 10

Deformation of Rocks





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Panamint Range, Death Valley National Park, California.

Chapter 10

Crustal Deformation

Topics covered -

- Stress and strain
- Mapping strike and dip
- Folds
- Faults
- Joints

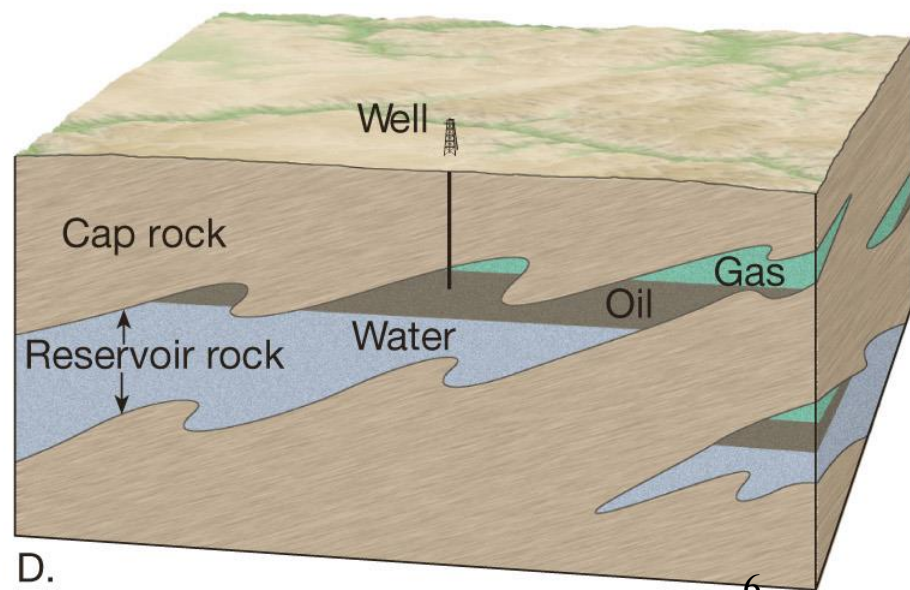
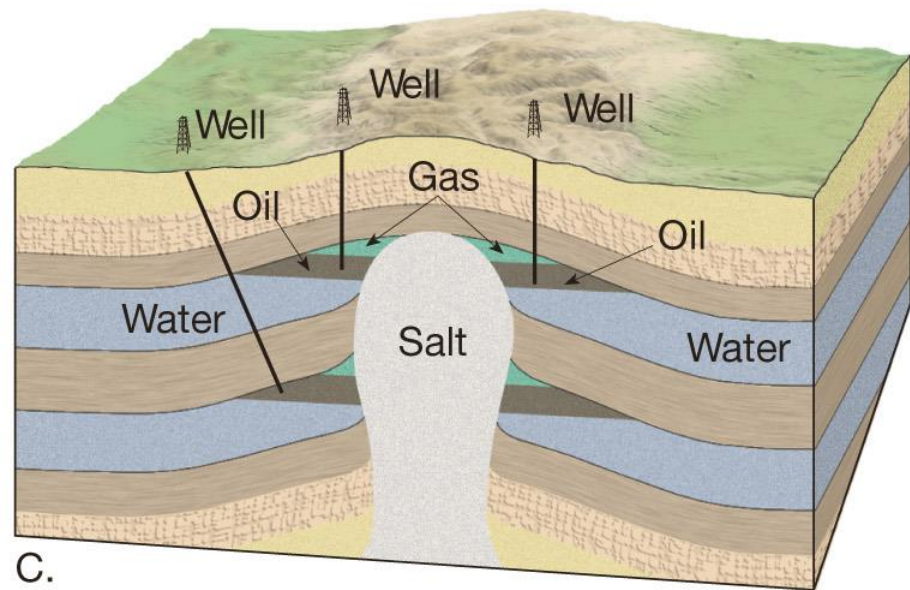
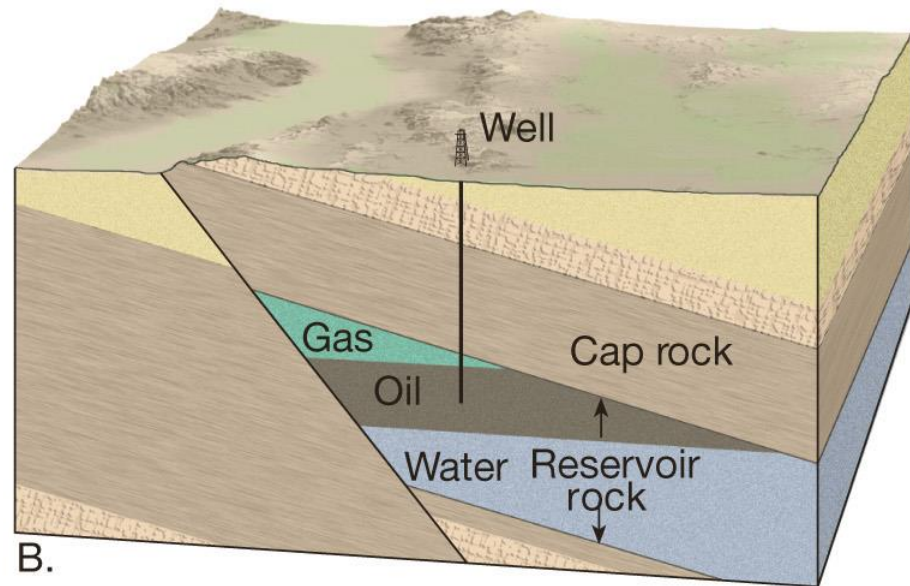
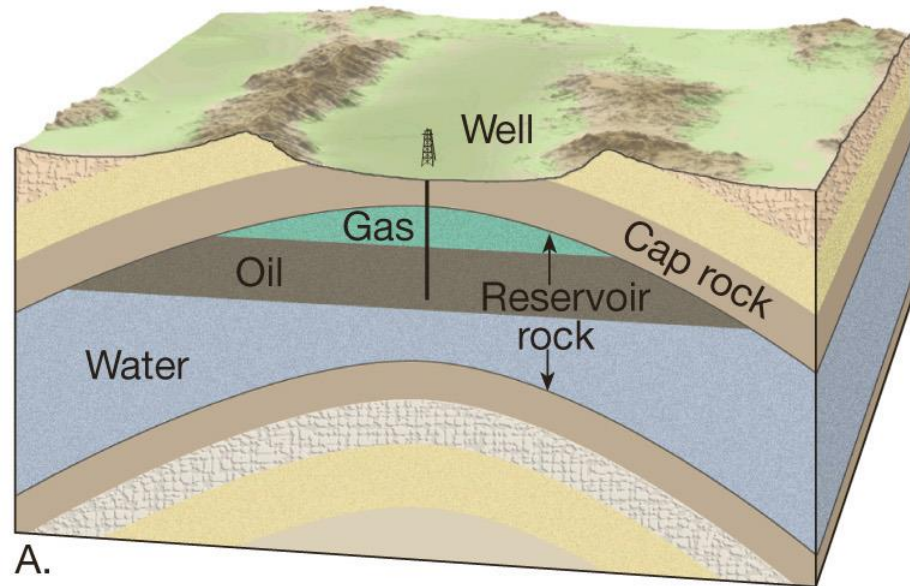
Why knowledge of folds and faults is important.

- orientation of folds and faults give
 - geologic setting
 - direction of forces that produces structures
- economic importance
 - oil and gas trapped by structures
 - must find faults so do not build structures there
 - ore deposits localized on faults

Why do oil companies spend a lot of money to find faults and folds?



Figure 21.7 Folds and faults control location of oil



Regional Growth Faults

Salt Dome

I-10

45

59

290

8

59

610

45

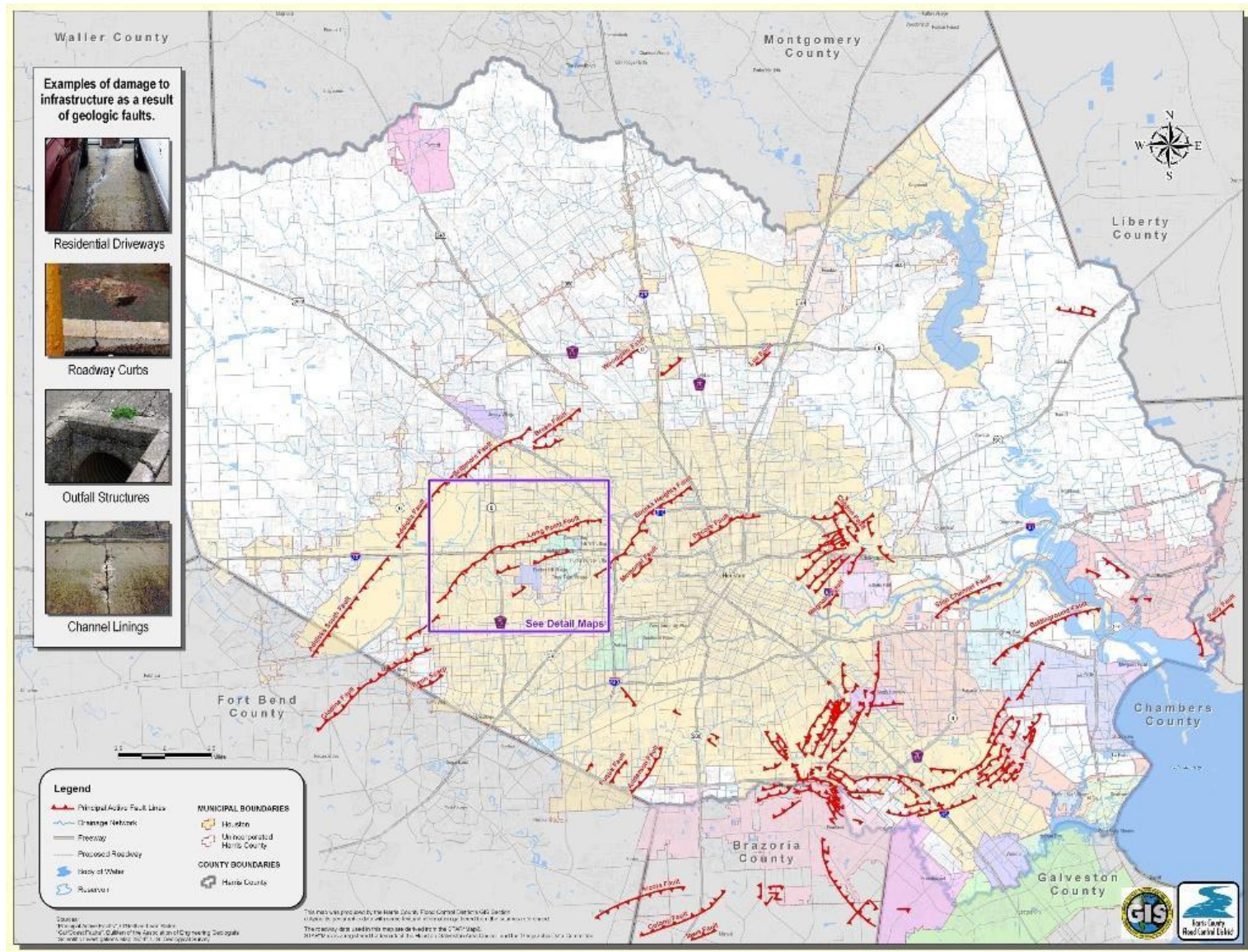
5 miles

N

The map displays a network of red lines representing faults, many of which are labeled with 'U' and 'D' to indicate upthrown and downthrown blocks. Yellow lines represent major highways, including I-10 running horizontally across the middle. Other roads are labeled with numbers: 45, 59, 290, 8, and 610. Two specific areas are circled in black, highlighting clusters of faults. A scale bar for 5 miles and a north arrow are located in the bottom left corner.

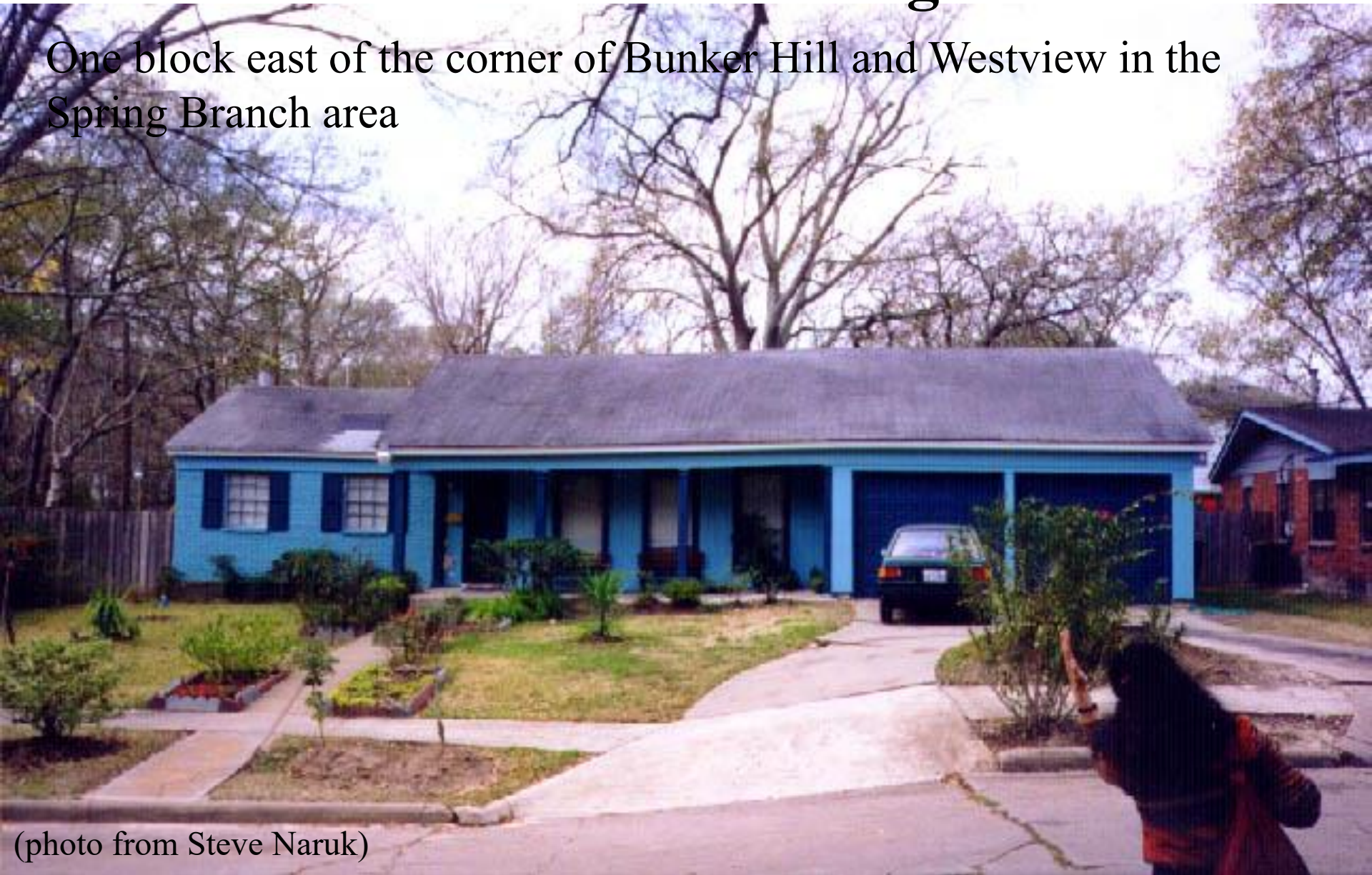
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Principal active faults in the Houston area (Harris County Flood Control District)



Faults in Houston can damage real estate

One block east of the corner of Bunker Hill and Westview in the Spring Branch area

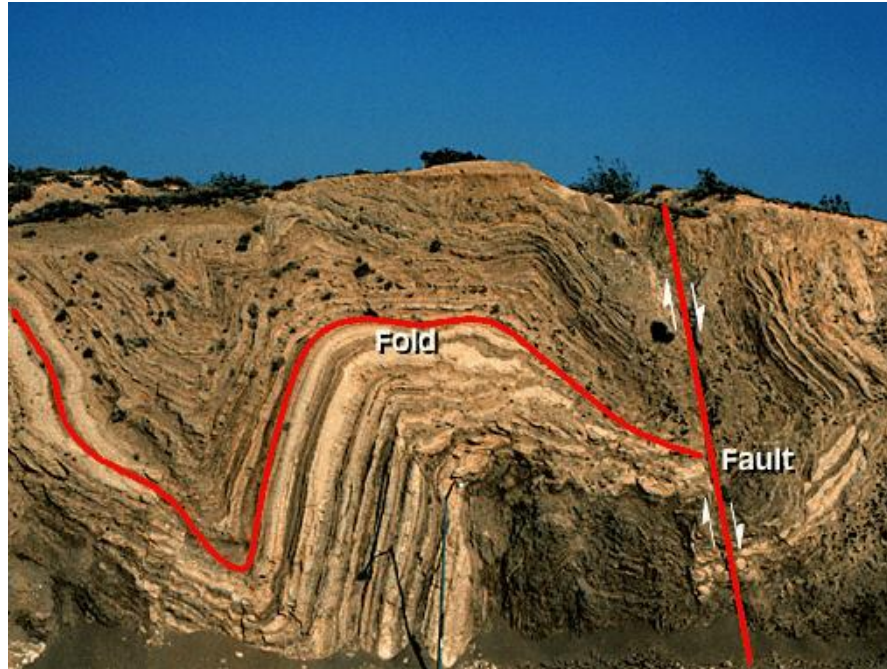


(photo from Steve Naruk)

Memorial Area, Houston, Texas

Deformation of Rocks

- Deformation results in folding, flowing, fracturing
- Folds and faults are geologic structures in response to tectonic forces or stress on the rock
- Structural geology is the study of the deformation of rocks.

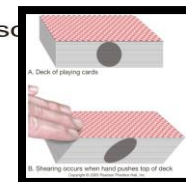
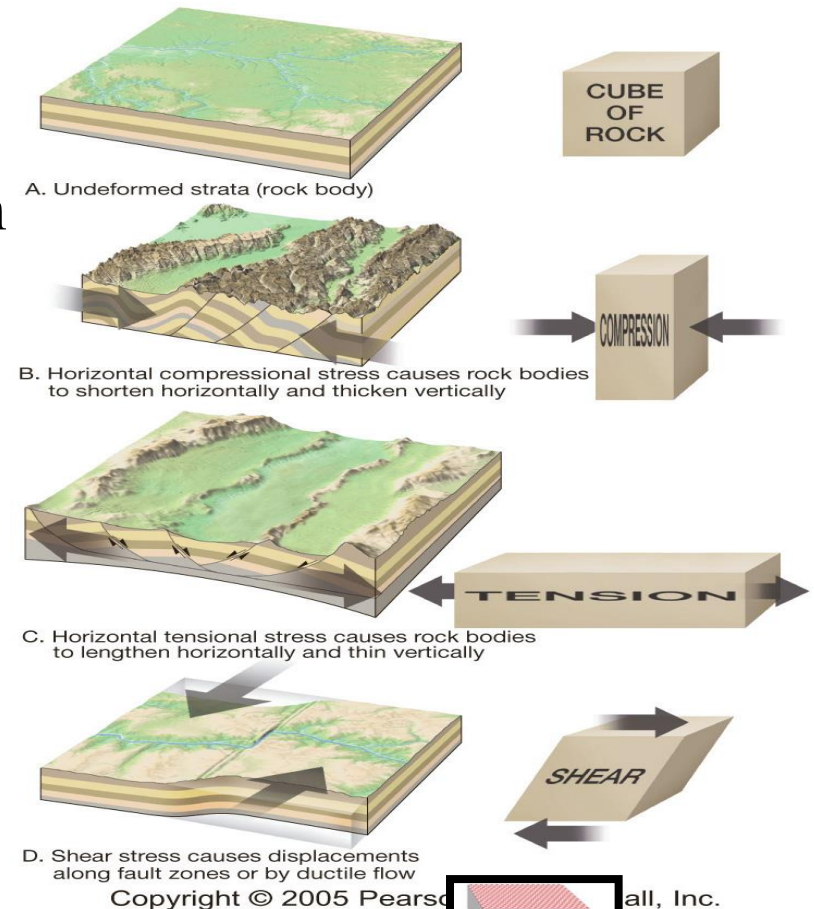


Stress and Strain

- Stress (tectonic forces)- is the amount of force acting on the rock
- Strain (deformation) – is the shape or volume changes caused by stress

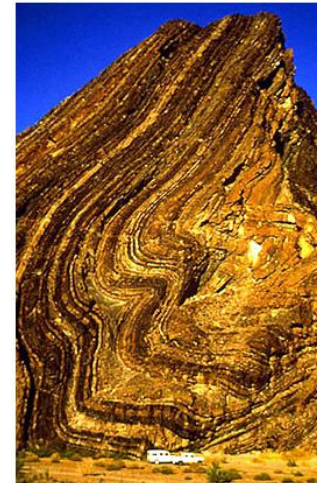
Three Types of Differential Stress

- Differential stress is force applied **unequally** in different directions.
- compressional stress – forces push together
 - shorten the rock body reducing its volume
- tensional stress (extensional) – forces pull apart
 - elongate rock body
- shear stress – forces push toward but parallel to one another



Three Types of Deformation

- elastic deformation -not permanent because the elastic limit of the rock has not been exceeded so the rock rebounds when the stress is released
- ductile deformation (ductile flow) - when the rock bends or flows, this is permanent.
- brittle deformation (brittle failure) – when the rock fractures (breaks), this is permanent



Factors that determine if ductile or brittle deformation will occur (that is will the rock fold or fracture)

- depth-
 - brittle deformation occurs at shallower depths because the rock is cooler and pressures are lower so the rock will break more easily.
 - ductile deformation occurs at great depth where pressure and temperature are higher.

Continued - Factors that determine if ductile or brittle deformation will occur

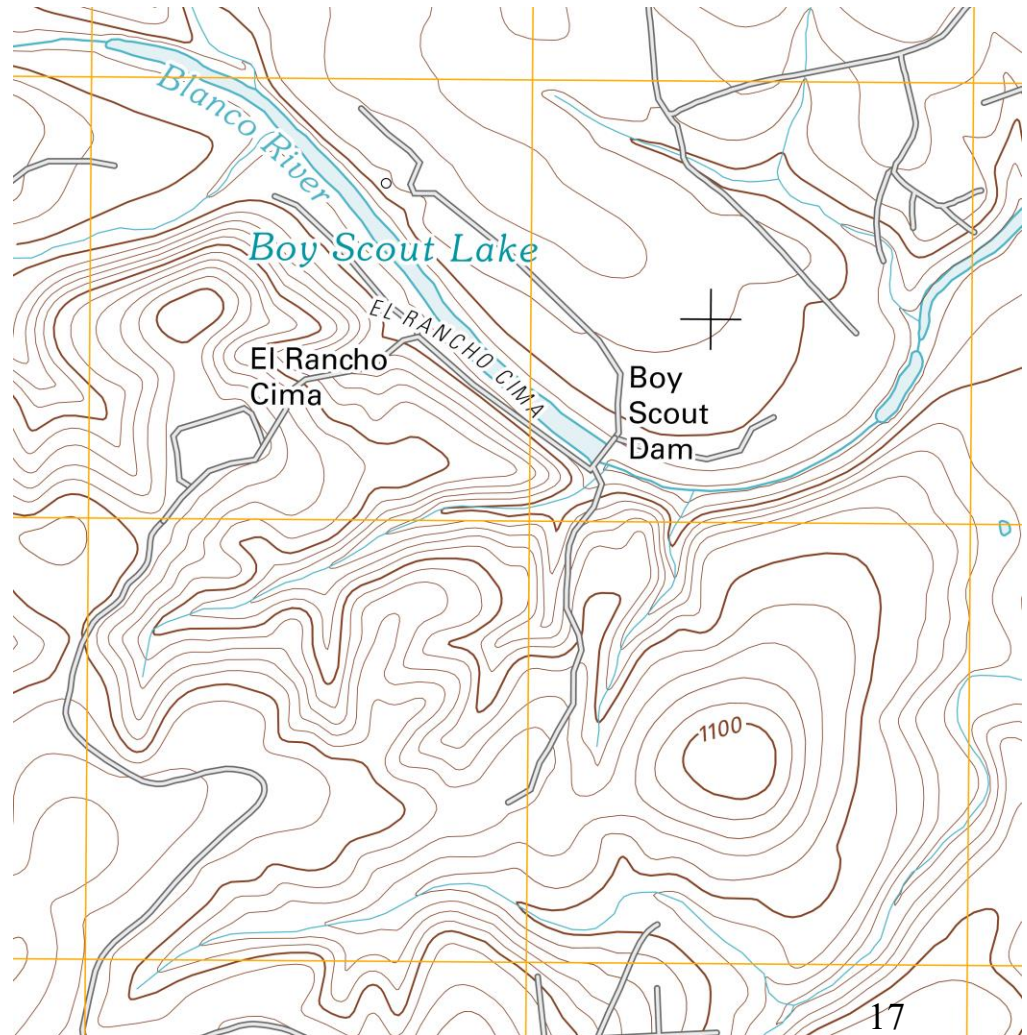
- rock type
 - brittle behavior is favored in granitic and basaltic rocks
 - Because they are composed of interconnected silicate minerals which have strong internal bonds
 - ductile behavior is favored in nonsilicate rich rocks or shale
 - because nonsilicate minerals have weaker bonds
 - because clays composing shale are weakly cemented
- time — brittle deformation occurs when forces acts quickly over short time, versus ductile deformation when forces acts slowly over long time

Mapping rock structure

- s • Locate outcrop - surface exposure of rock
 - Identify the rock type
- Measure strike and dip of folds, faults and rock layers
- Plot all the strike, dip and rock type information on a geologic map.

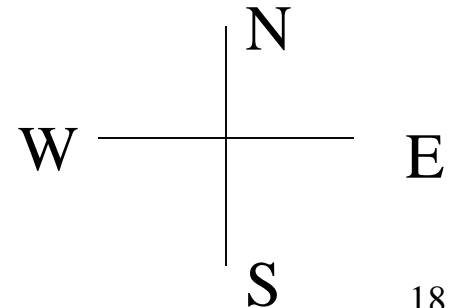
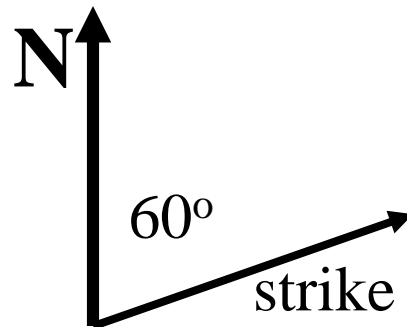
Topographic maps used as basemap

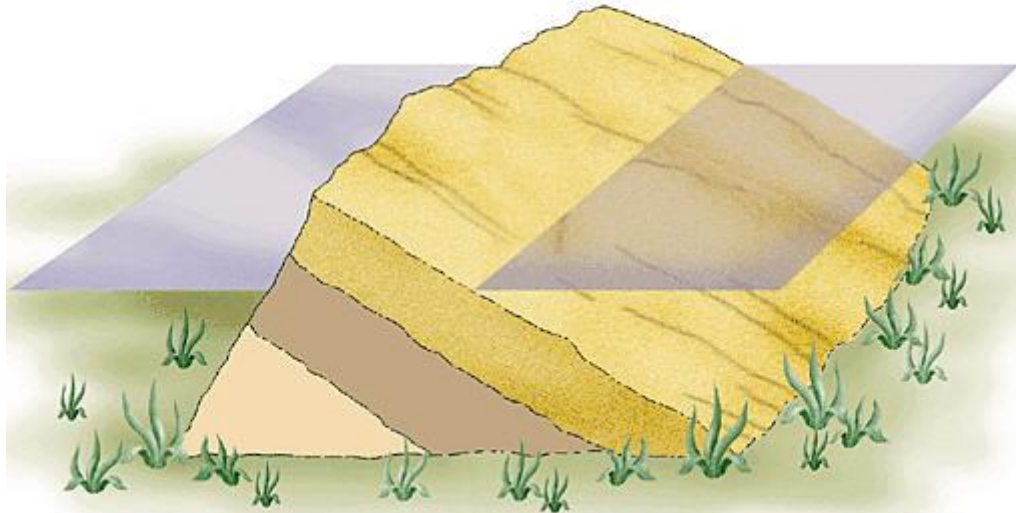
- Contour lines of constant elevation.
- Free Topographic Maps for all of the US can be downloaded from the United States Geological Survey web site
[http://store.usgs.gov/b2c_usgs/usgs/maplocator/\(xcm=r3standardpitrex_prd&layo=ut=6_1_61_76&uiarea=2&ctype=areaDetails&carea=%24ROOT\)/.do](http://store.usgs.gov/b2c_usgs/usgs/maplocator/(xcm=r3standardpitrex_prd&layo=ut=6_1_61_76&uiarea=2&ctype=areaDetails&carea=%24ROOT)/.do)



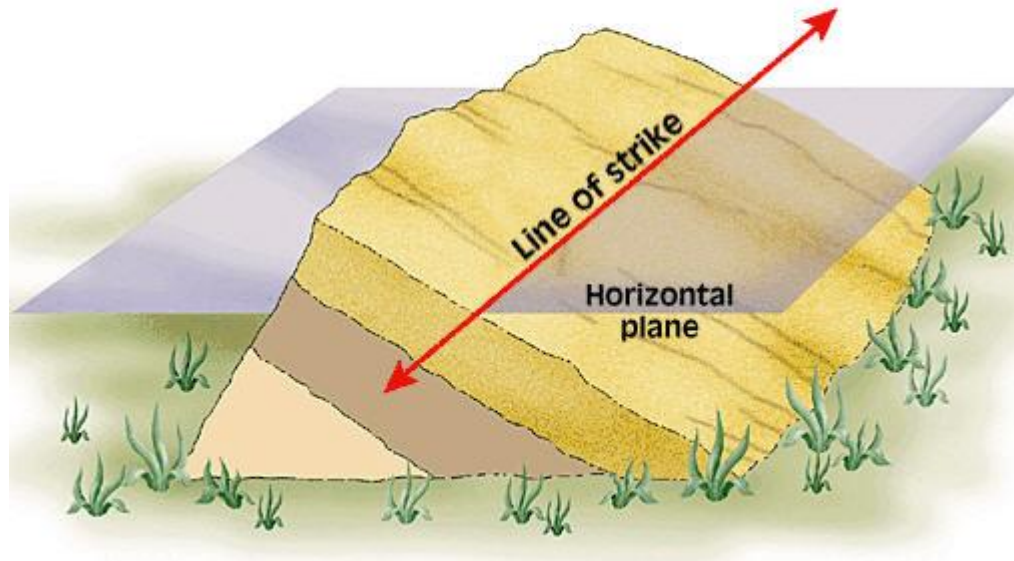
Strike

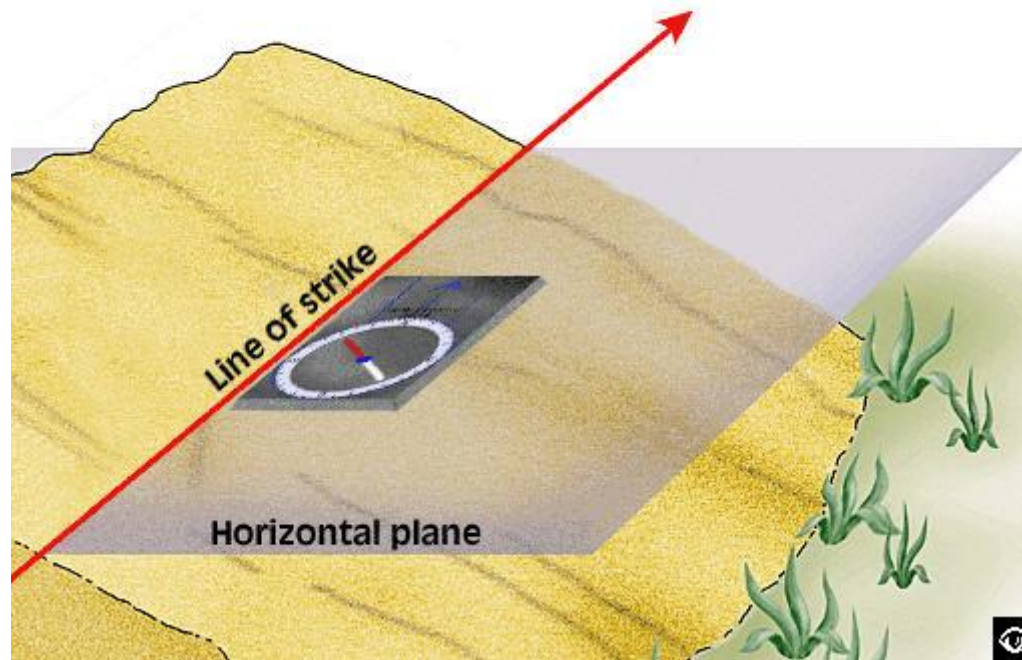
- Line produced by the intersection of a horizontal plane on the rock surface
 - if roll ball down rock surface it is the line perpendicular to the path the ball takes
- report as an angle relative to north
 - N60°E = strike is 60° east of north

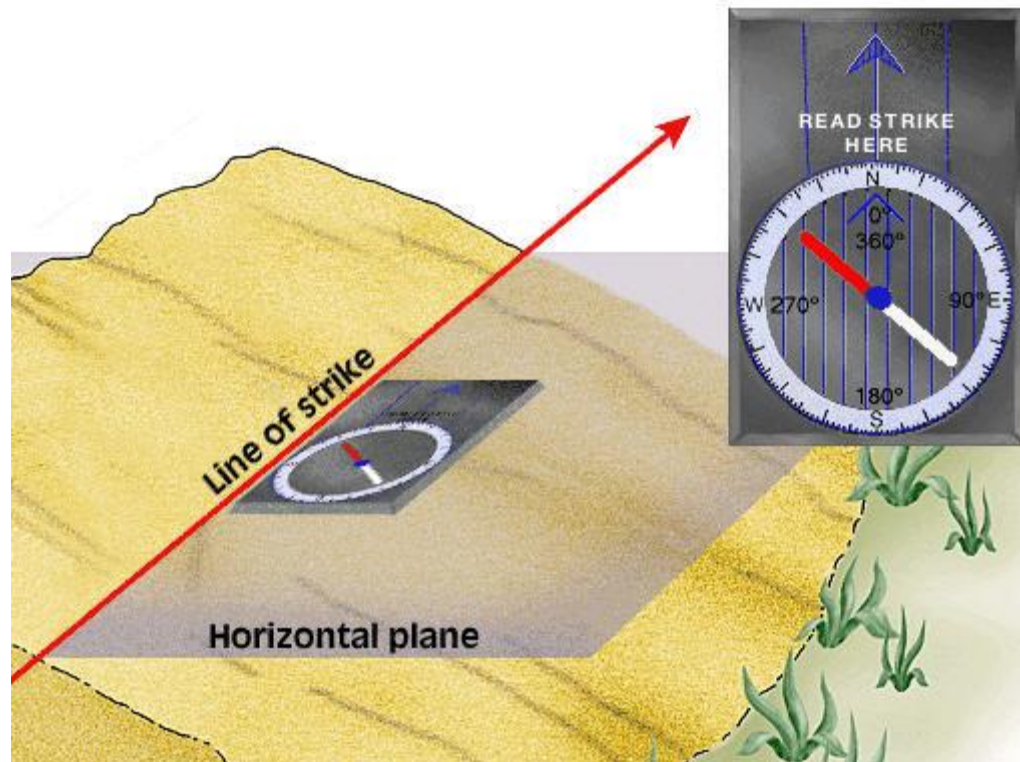


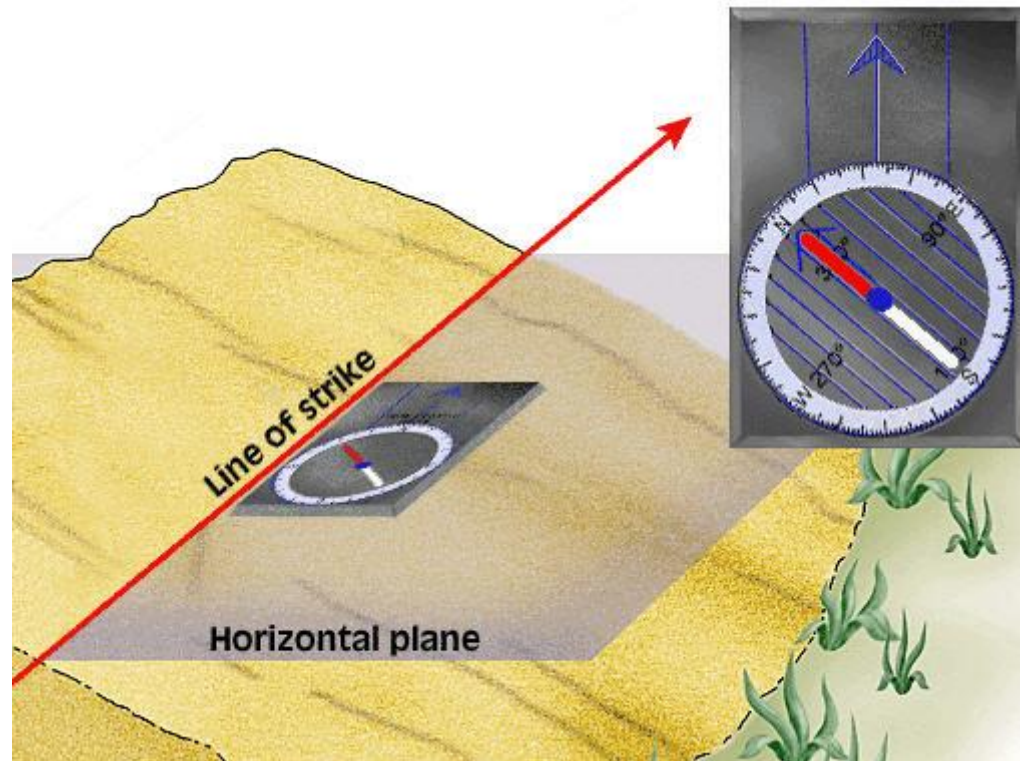


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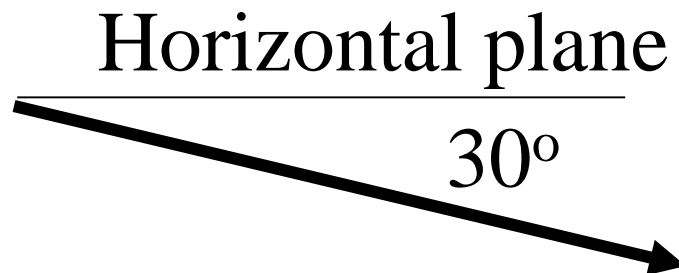


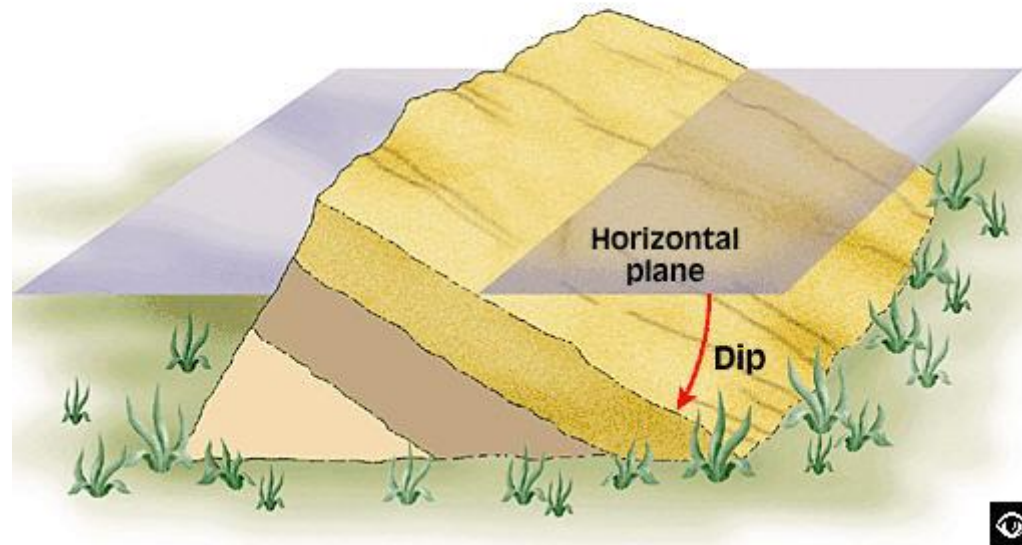




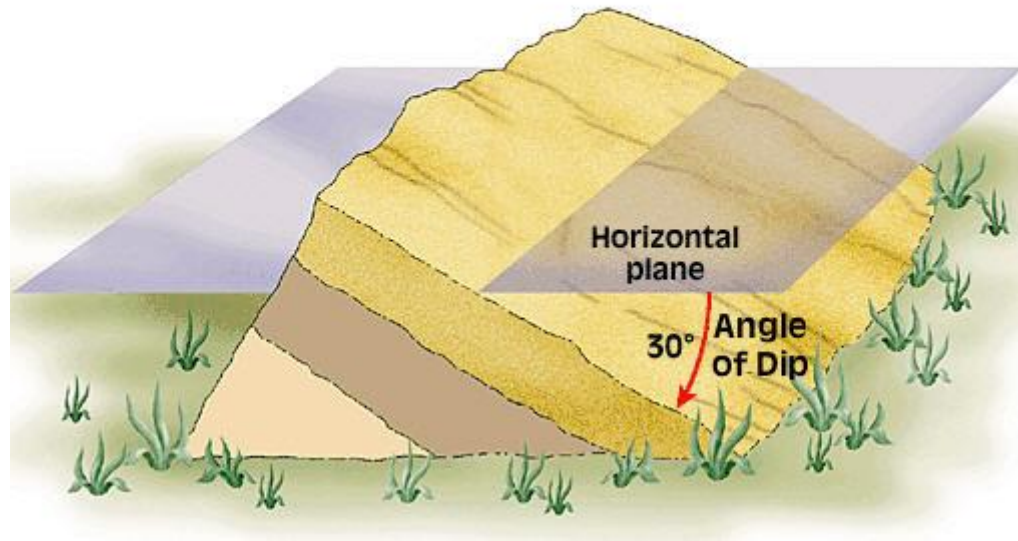
Dip

- angle of inclination of the rock surface from the horizontal plane
 - the path of a ball rolling down the rock face
- dip is always perpendicular to the strike
- report angle and direction of dip
- dip of 30°E

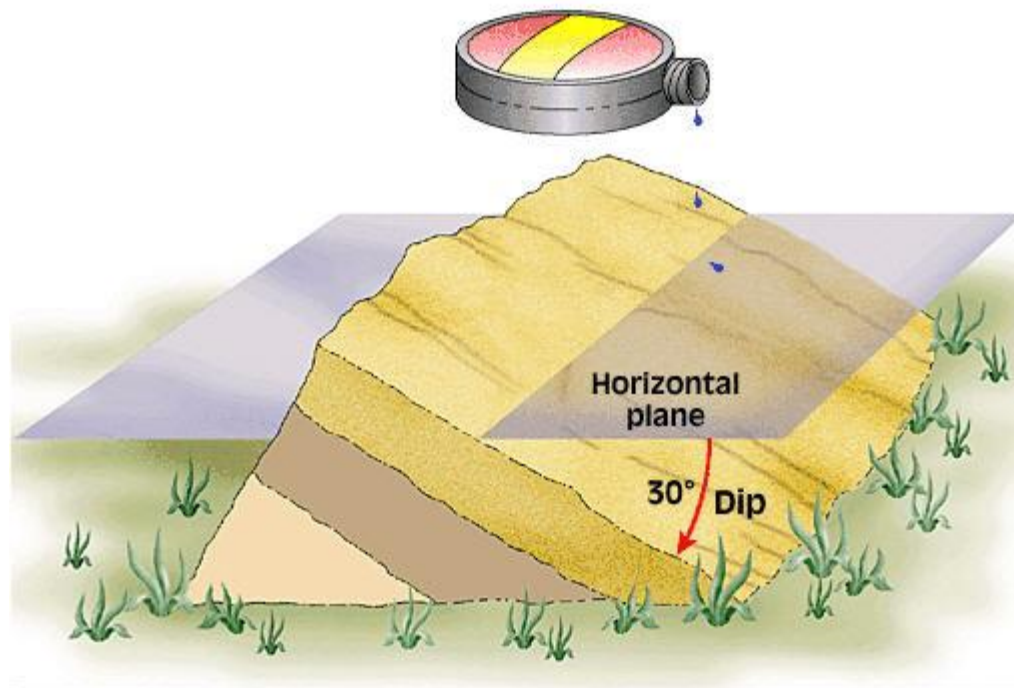


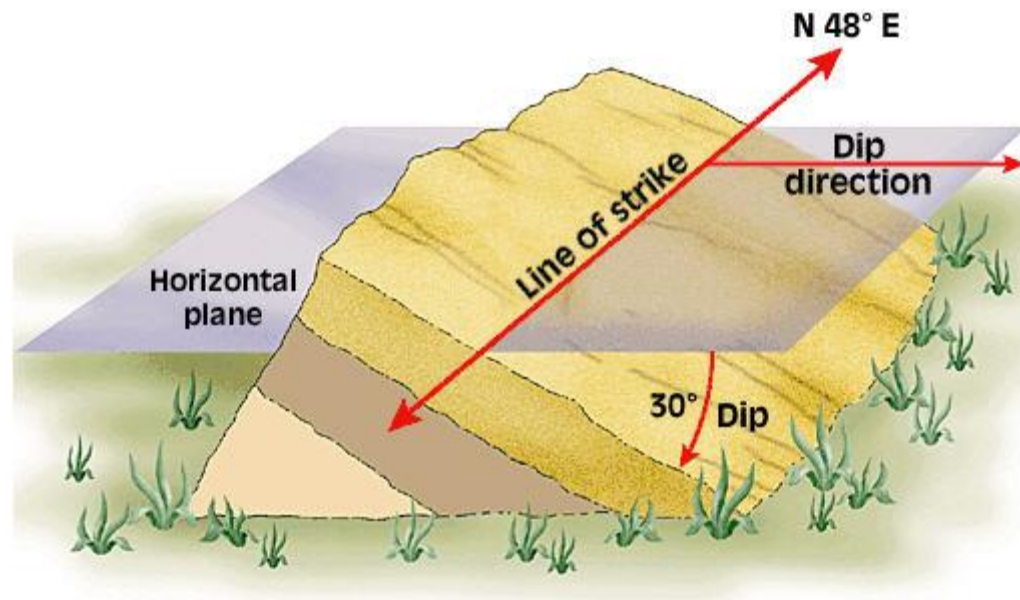


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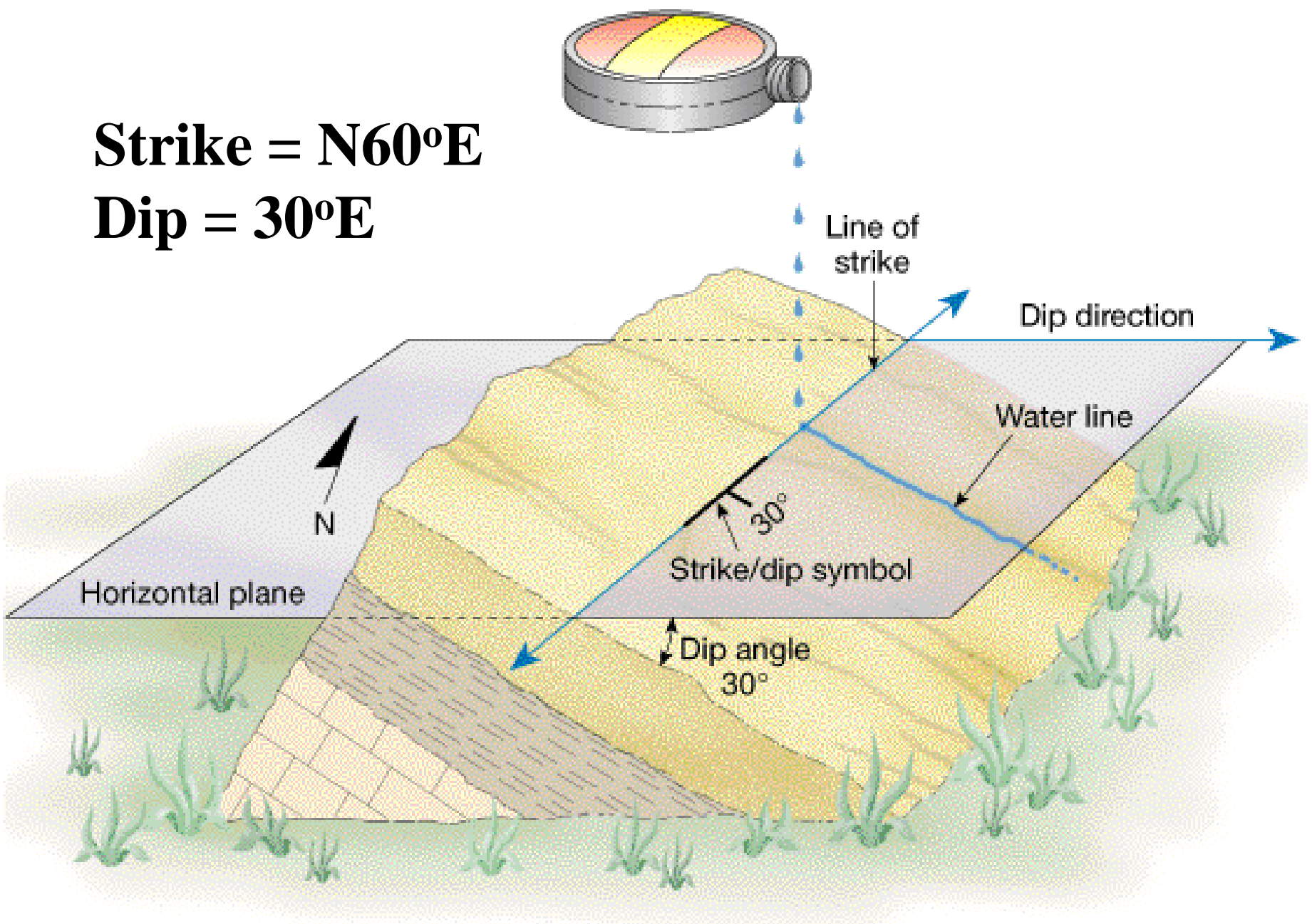
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Strike = N60°E

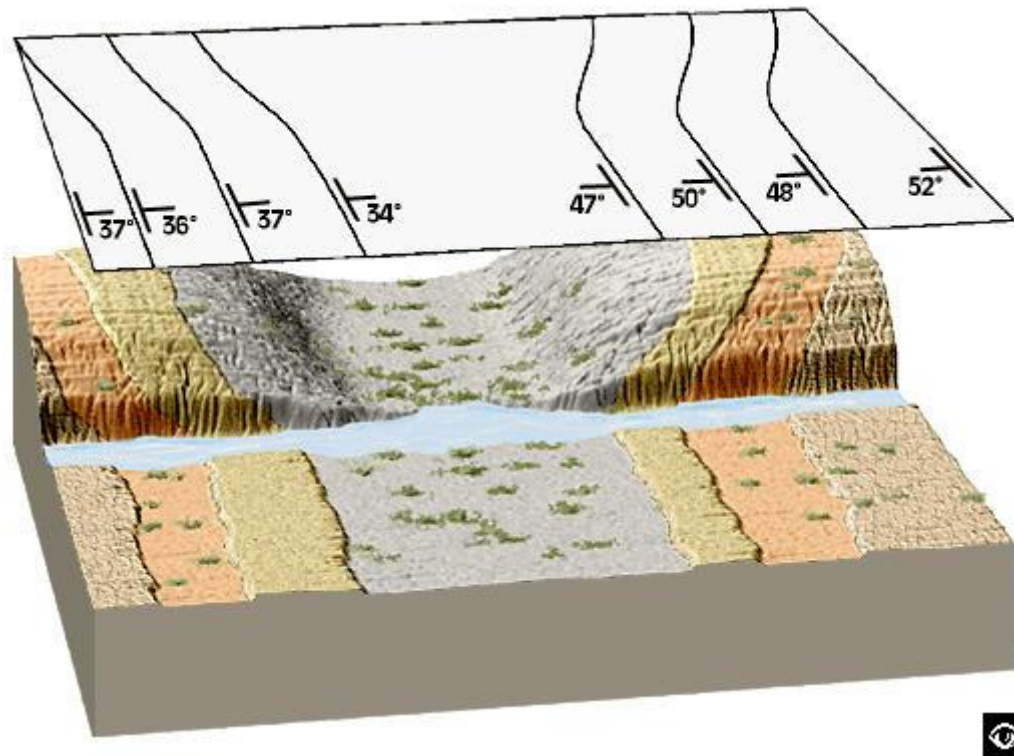
Dip = 30°E



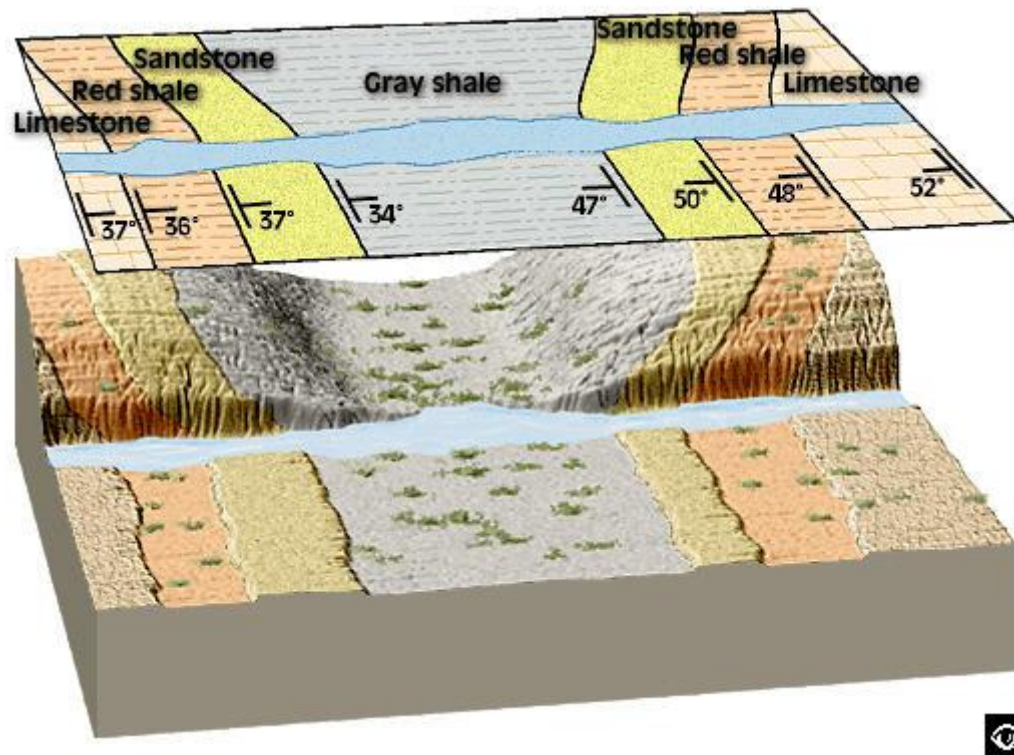
Geologic map

- plot strike and dip of each structure
- color code rock formations
- can use this map to infer the shape and orientation of structure below the ground

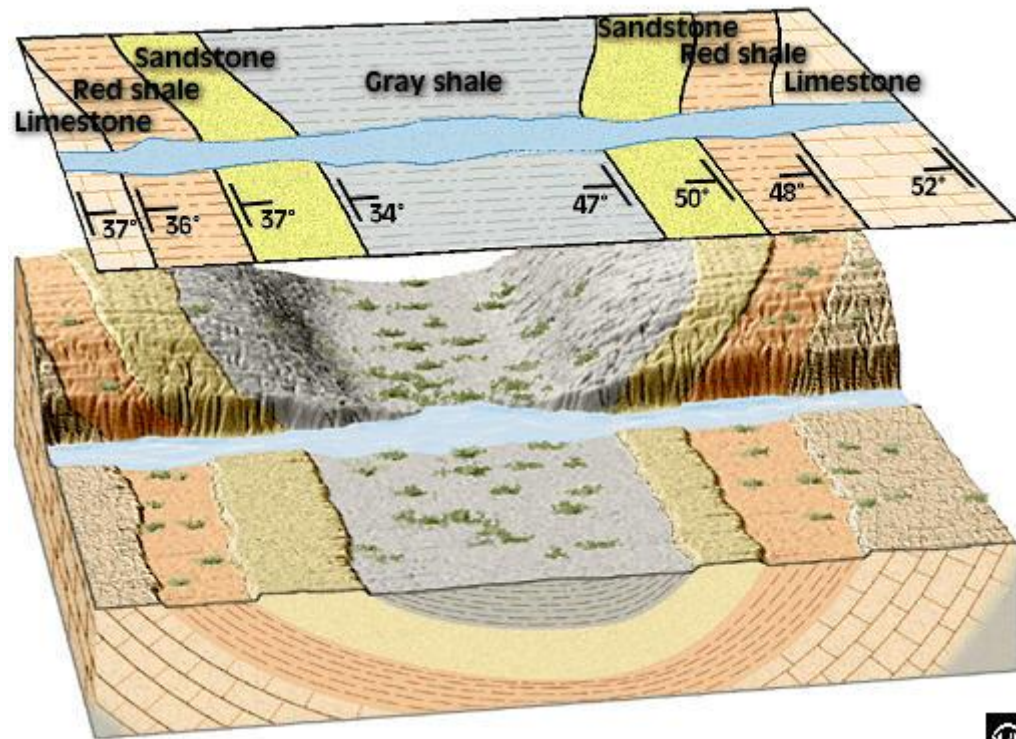
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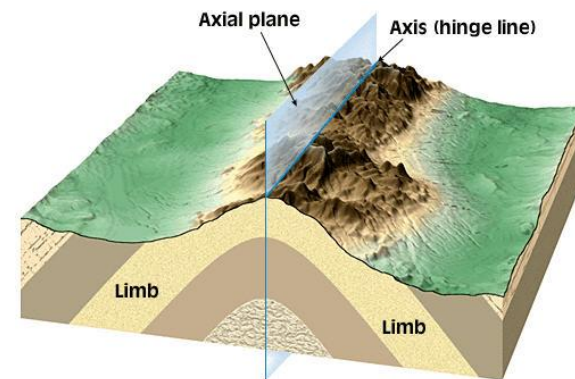
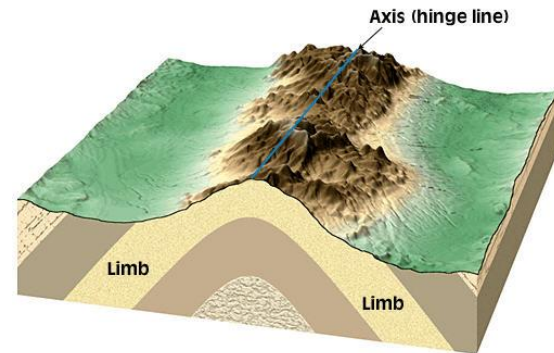
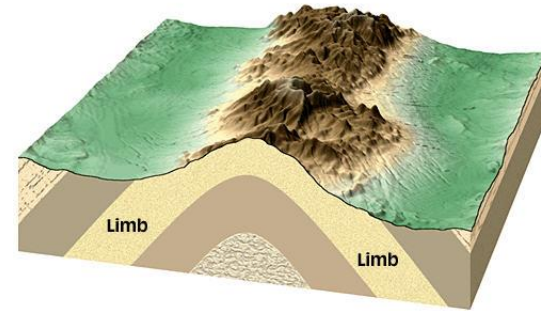
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FOLDS

Parts of a Folded Rock

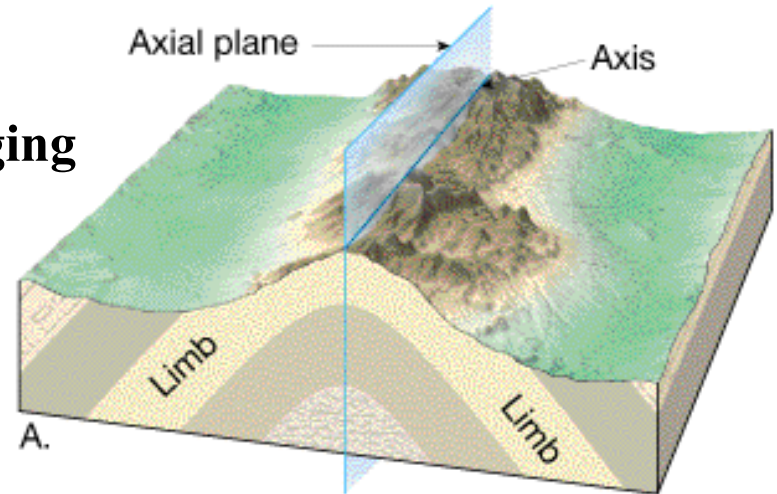
- limb - the 2 sides of the fold
- hinge line (axis) - a line drawn along the points of maximum curvature of each layer
- axial plane - imaginary surface that divides the fold as symmetrically as possible



Parts of a Folded Rock (continued)

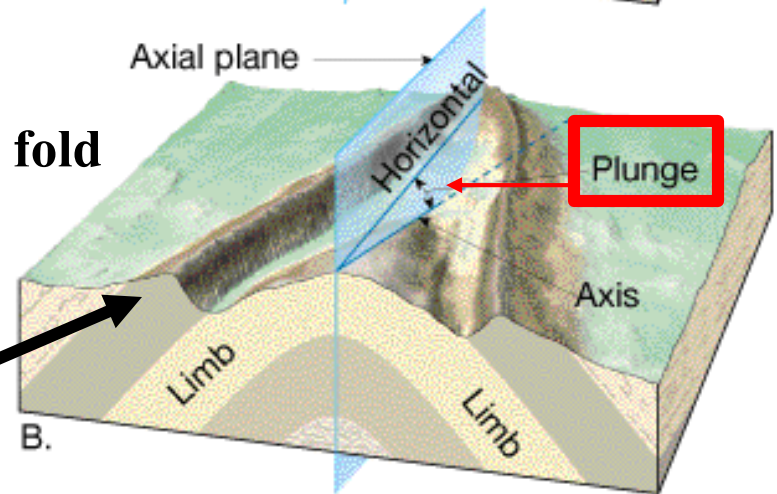
- plunge - the angle of incline of the axis
- hogback - steeply inclined angular ridges
 - formed when folded strata are resistant to weathering and form outcrop

Non-plunging fold



Plunging fold

hogback

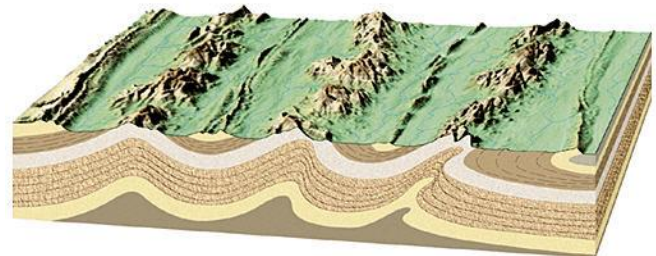


Fold types

- anticline/syncline
- monocline
- dome/basin

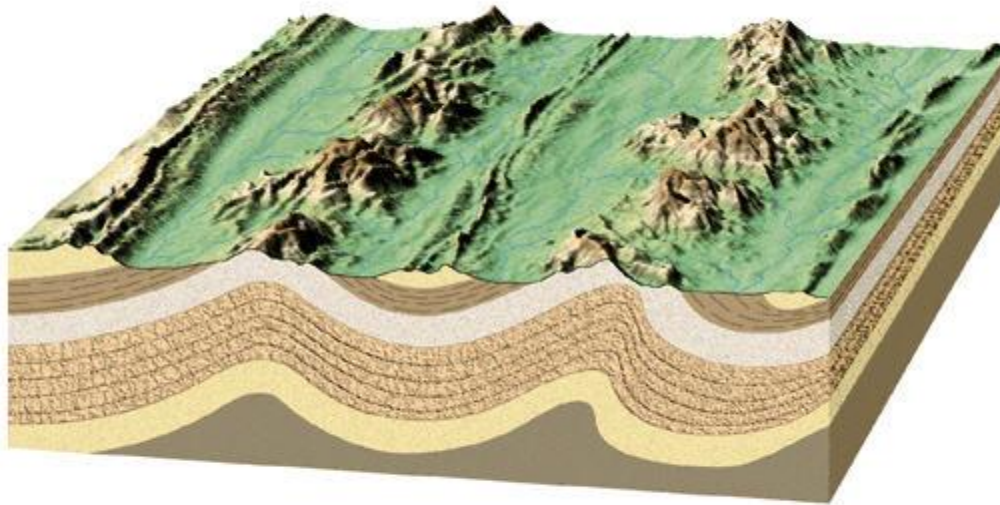
Anticline and Syncline

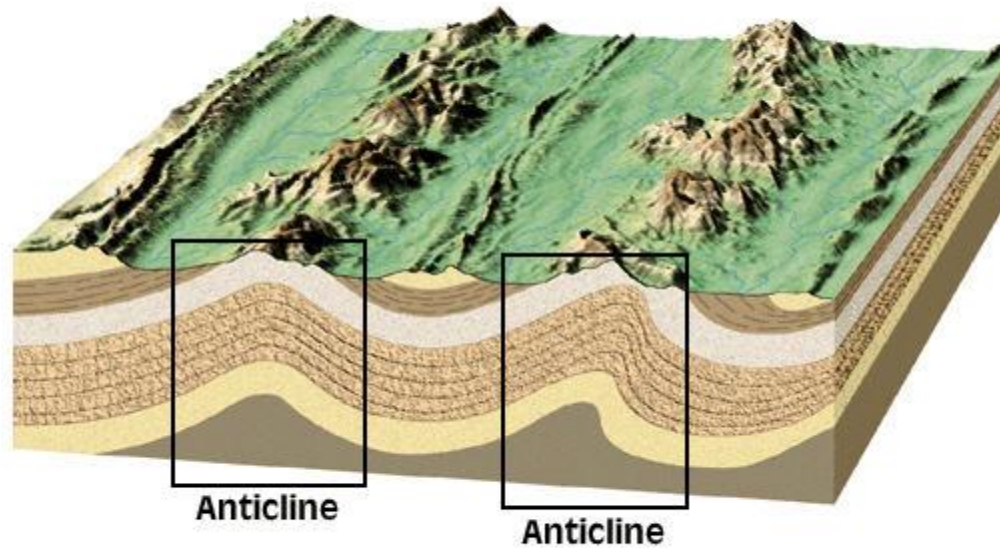
- anticline - arching of geologic layers
 - older rocks on the inside
- syncline - trough, downfolded layers
 - older rocks on the outside
- produced by compressional forces
- occur together in sets



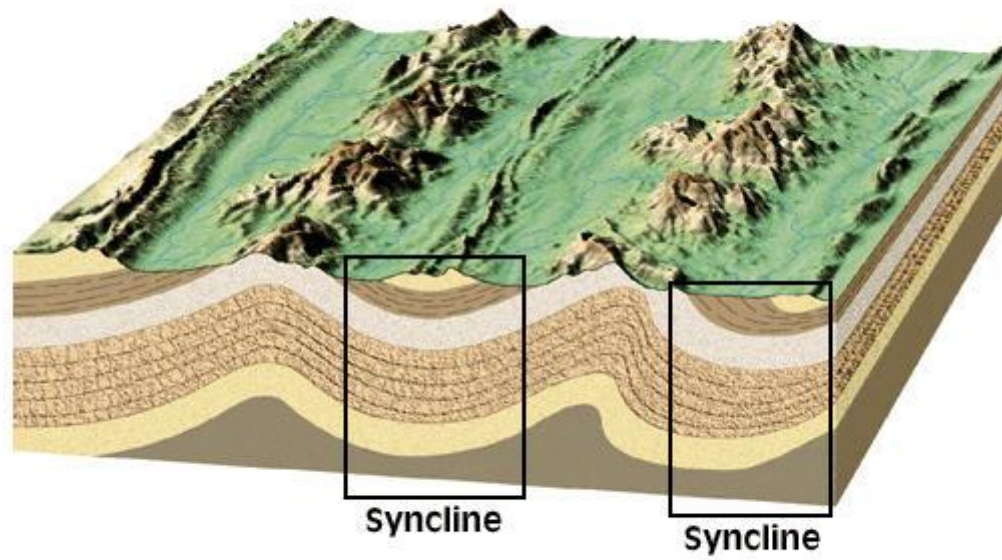
Features of anticlines and synclines

- symmetrical when both limbs are at same angle
- asymmetrical when the limbs are at different angle
- plunging when the hinge line dips into the ground
 - anticline outcrop points in the direction of plunge
 - syncline outcrop points opposite of the plunge direction
- these folds do not go on forever they dissipate at their ends

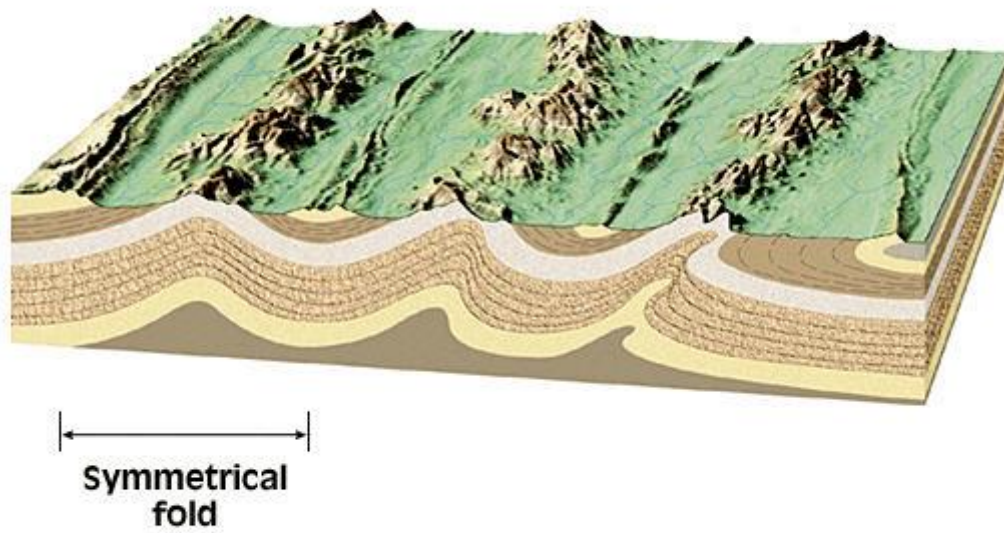




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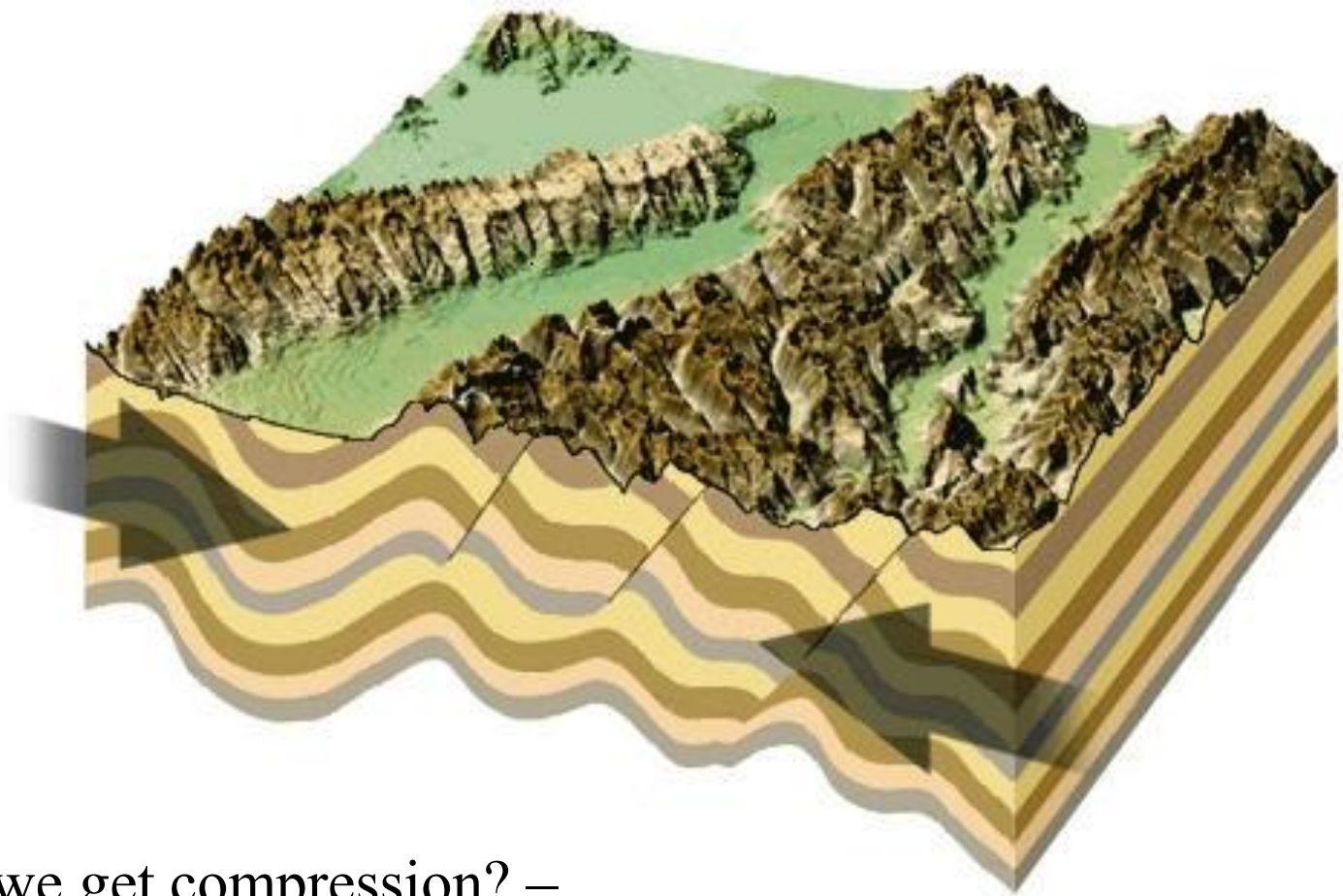


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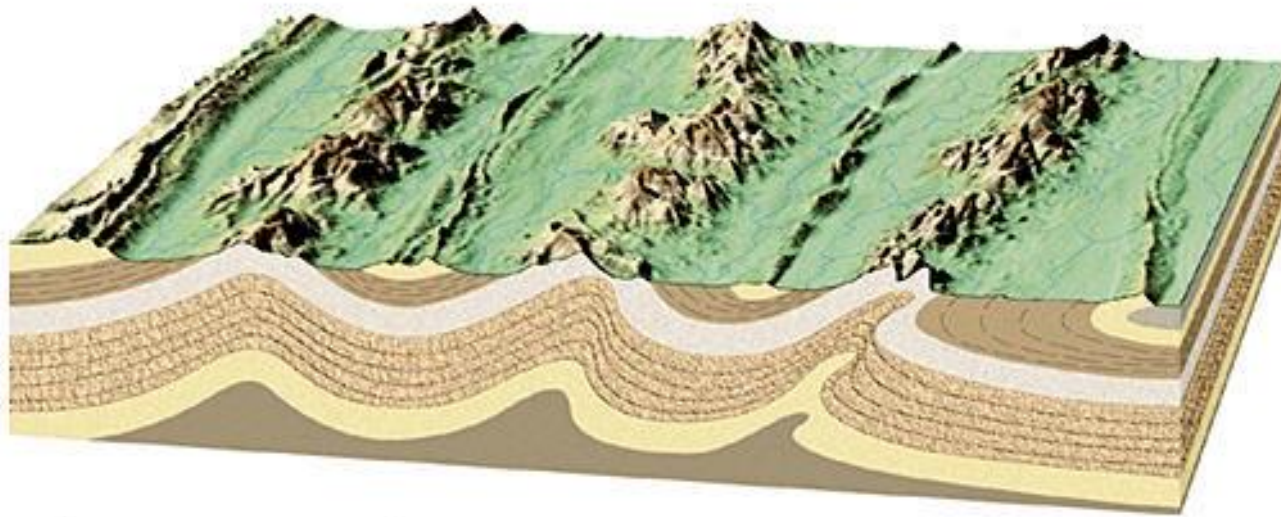
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Compressional forces produce anticlines and synclines

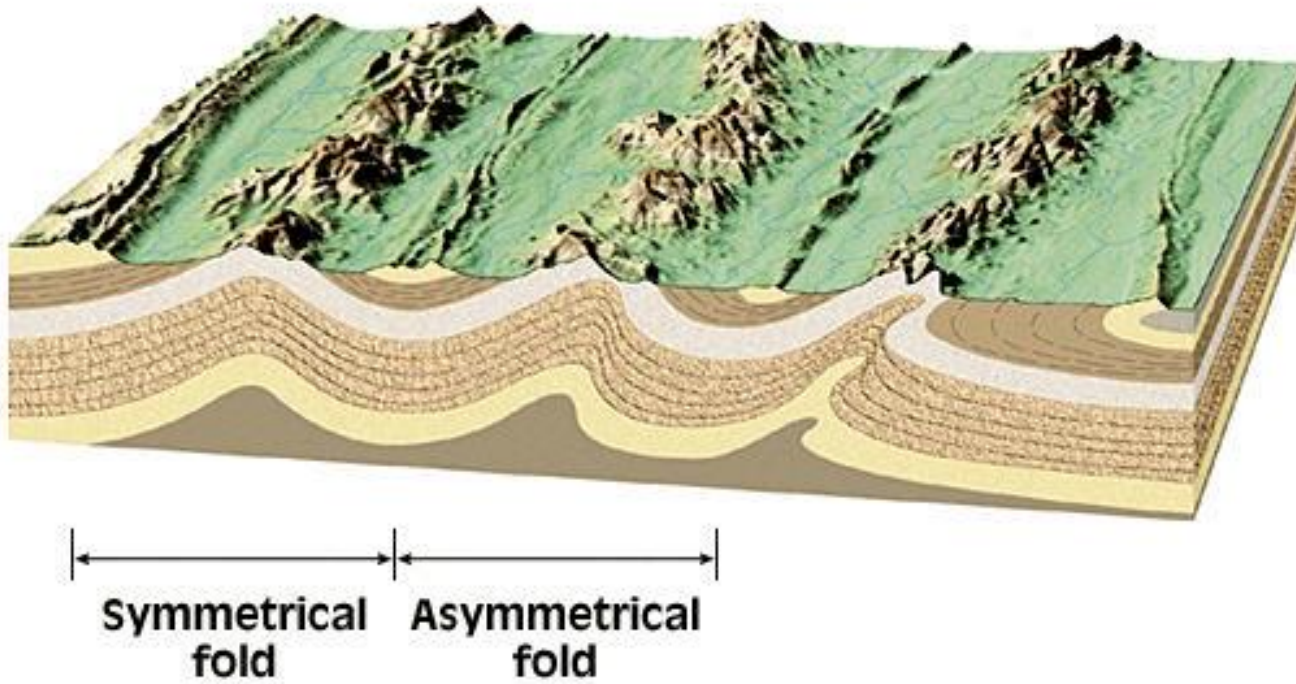


Where do we get compression? –
Subduction zones where plates are colliding





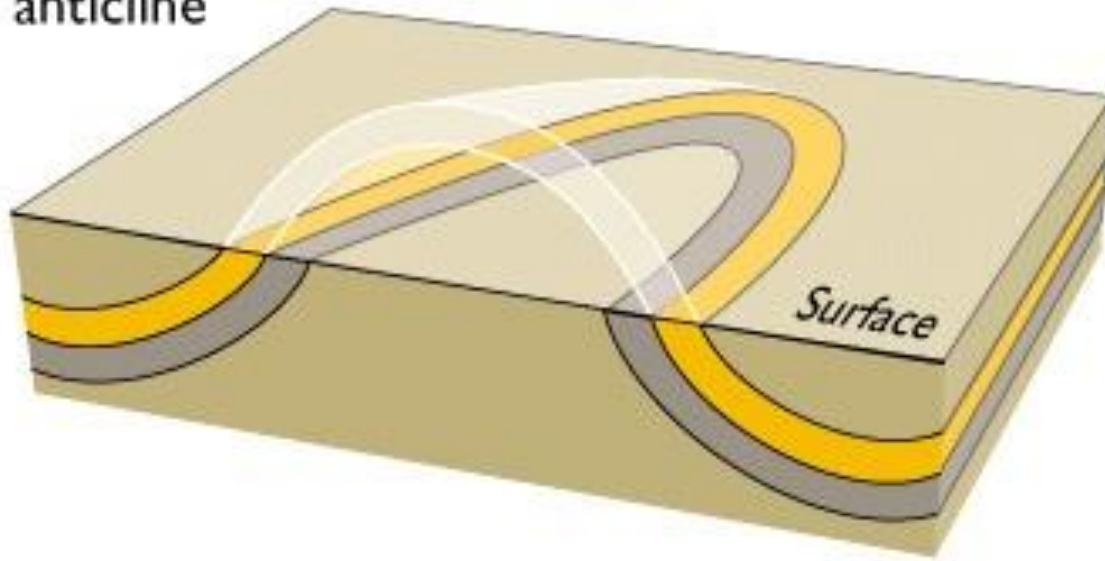
←→
**Symmetrical
fold**



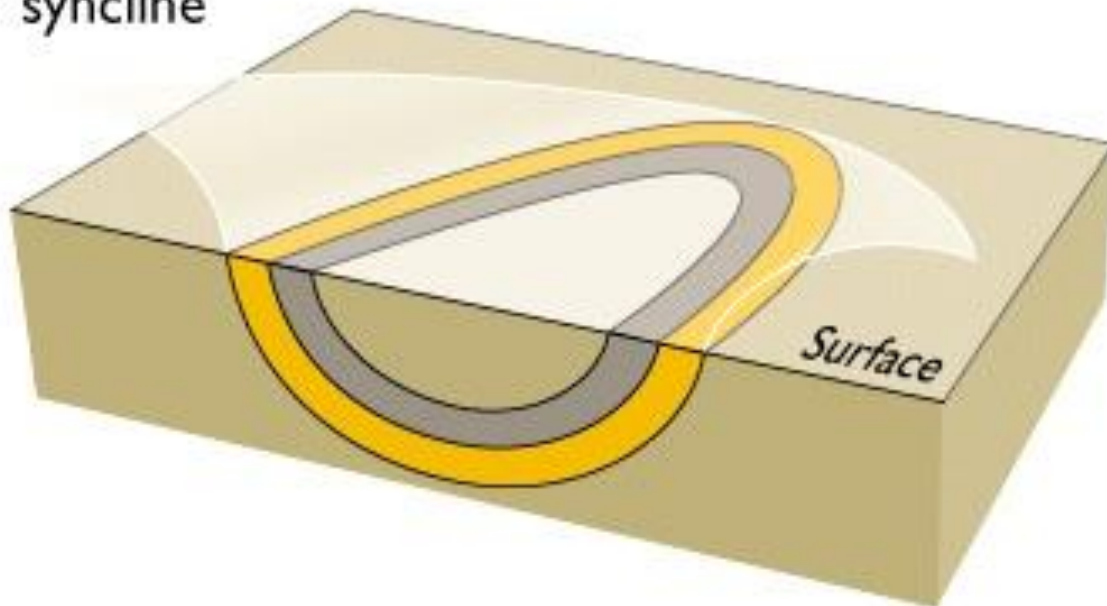
Asymmetric Folds



Plunging
anticline

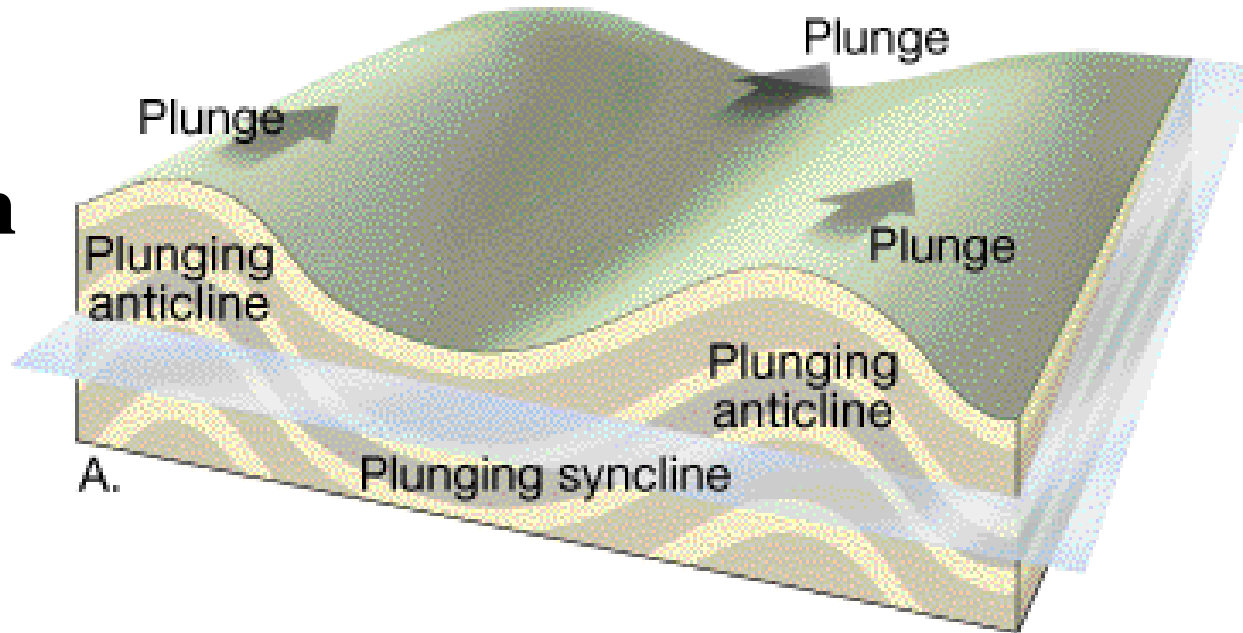


Plunging
syncline

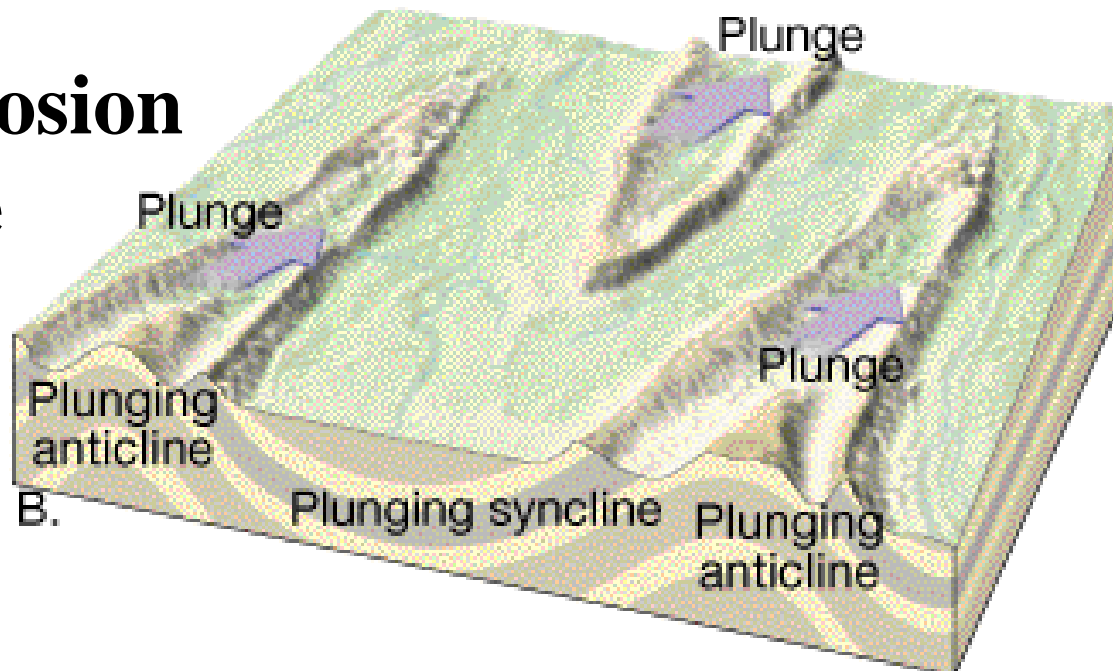


Map View of Plunging Folds

**Before
erosion**



**After erosion
of to the
plane
shown
above**



An aerial photograph of a desert landscape with a prominent geological feature. A dashed white line is drawn across the image, following the crest of a plunging anticline. The terrain is arid, with various shades of brown, tan, and green. In the lower-left quadrant, there is an oil field with several structures, including storage tanks and buildings. The overall scene illustrates the relationship between the geological structure and the oil field's location.

Axial Trace of a Plunging Anticline

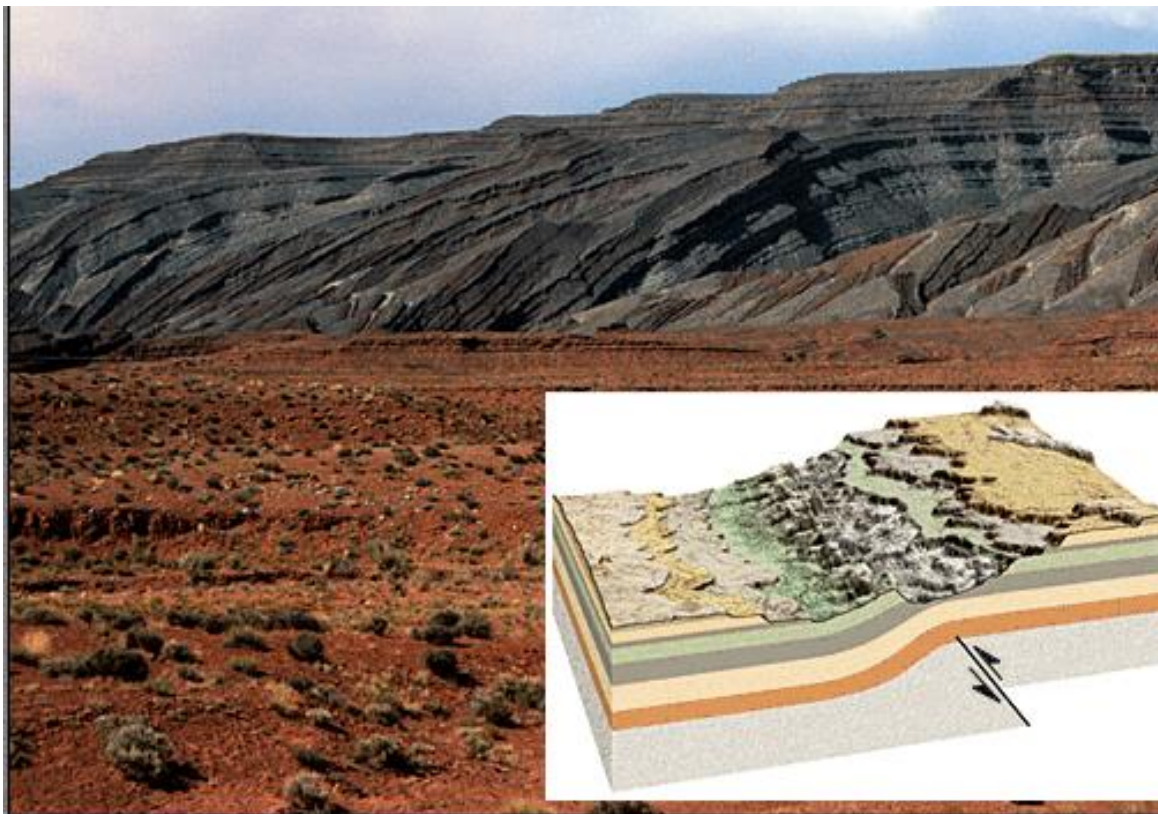
Landers Oil Field which
occurs
on the crest of an anticline



C. Sheep Mountain Anticline, Wyoming, a plunging anticline.

Monocline

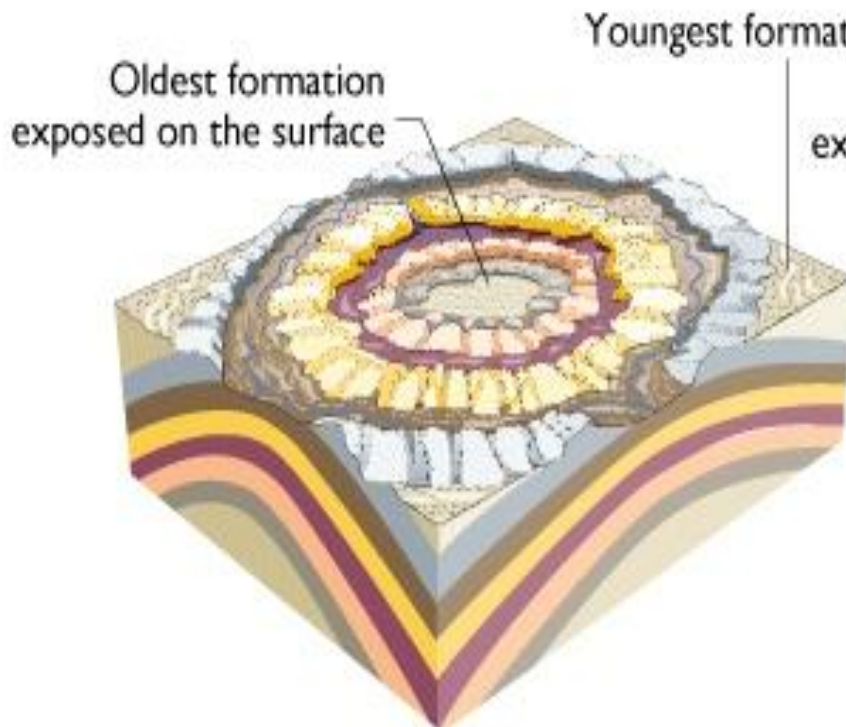
- A fold that has only one limb
- produced by vertical displacement along a fault rather than compression
- Overlying folded strata are folded over a vertical fault in the underlying basement rock



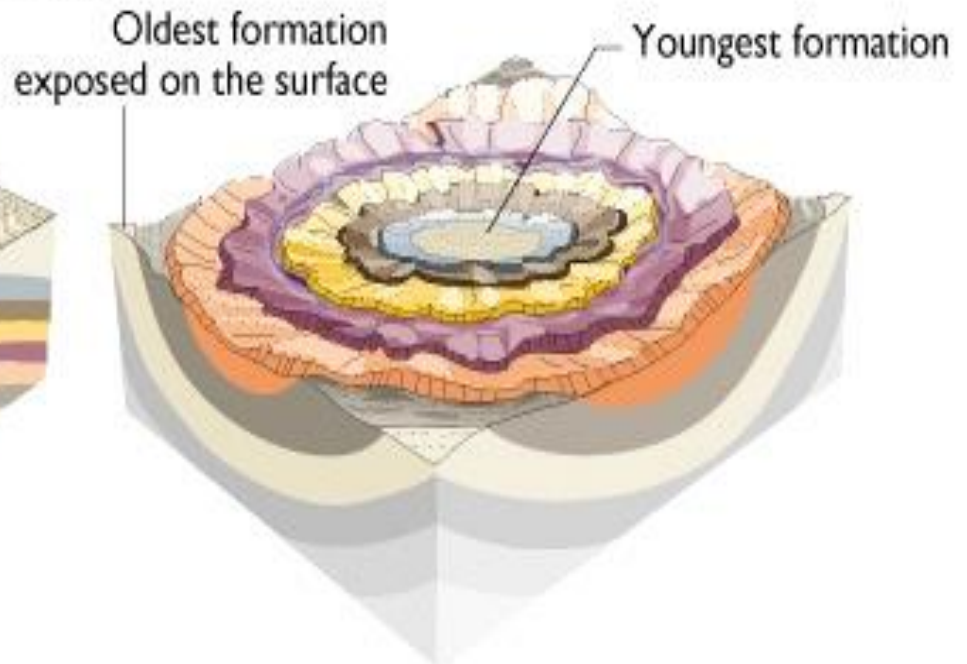
Dome and Basin

- dome - circular upwarping, oldest strata inside
- basin - circular downwarping, youngest strata inside

Dome



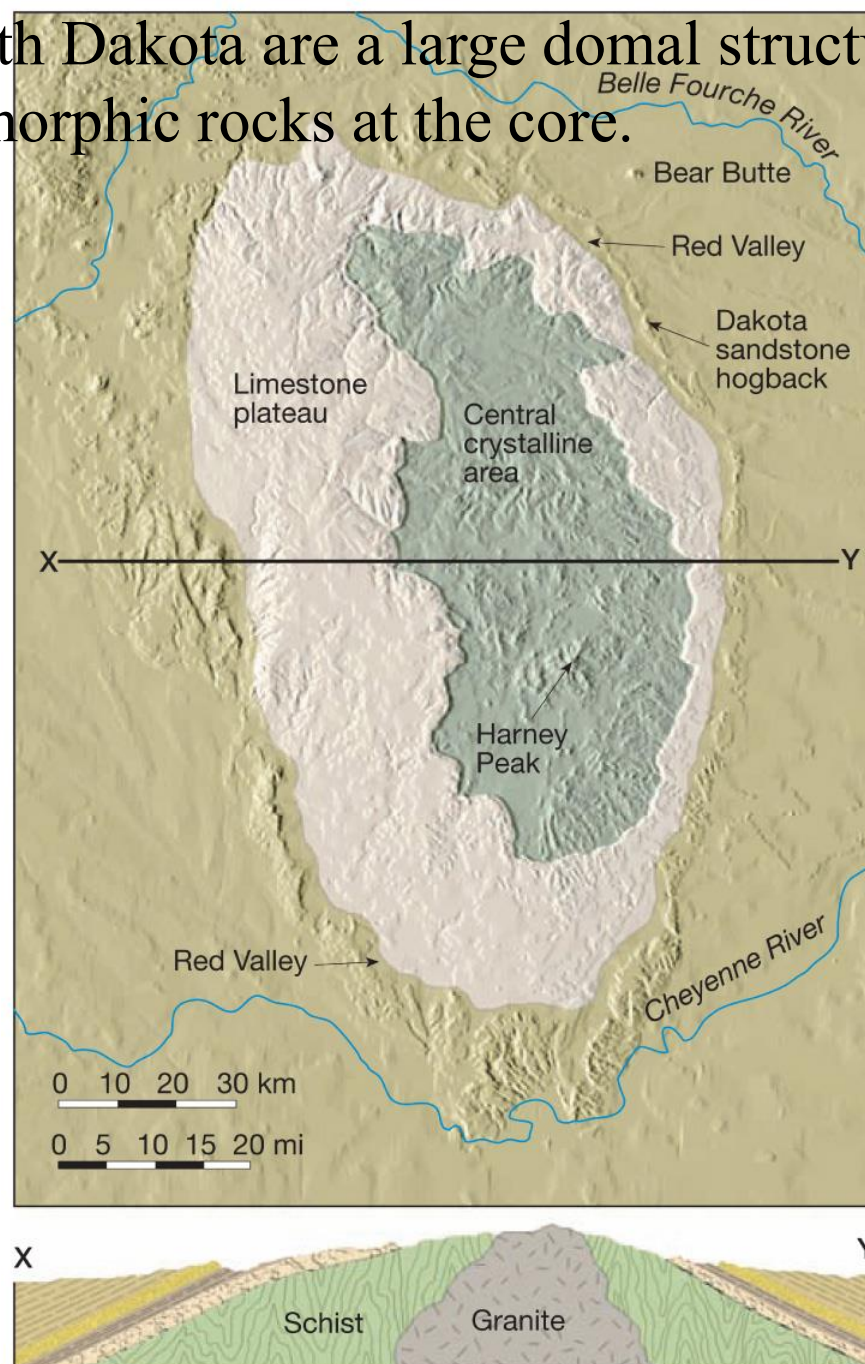
Basin



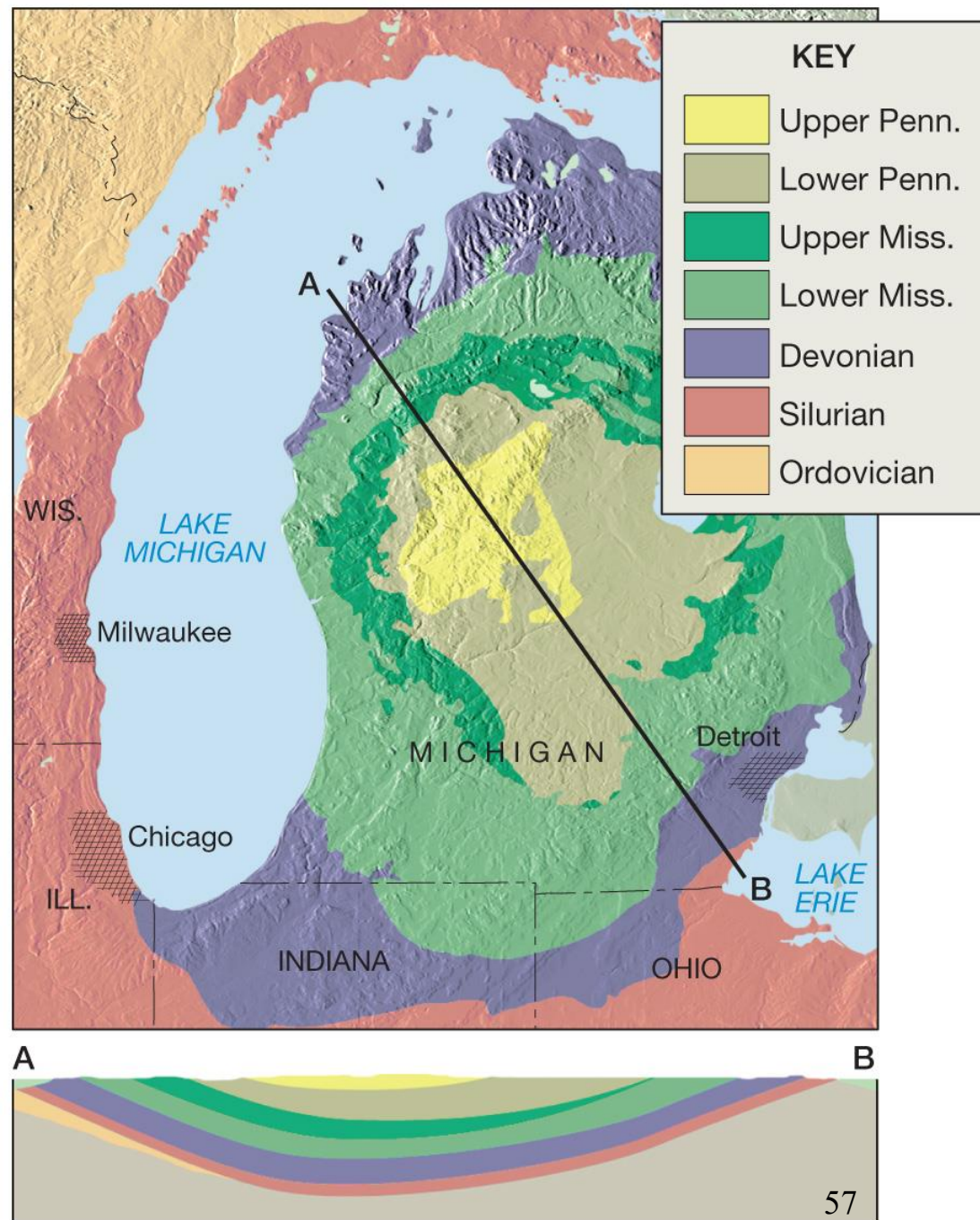
Sinclair Dome, Wyoming



Black Hills of South Dakota are a large domal structure with resistant igneous and metamorphic rocks at the core.



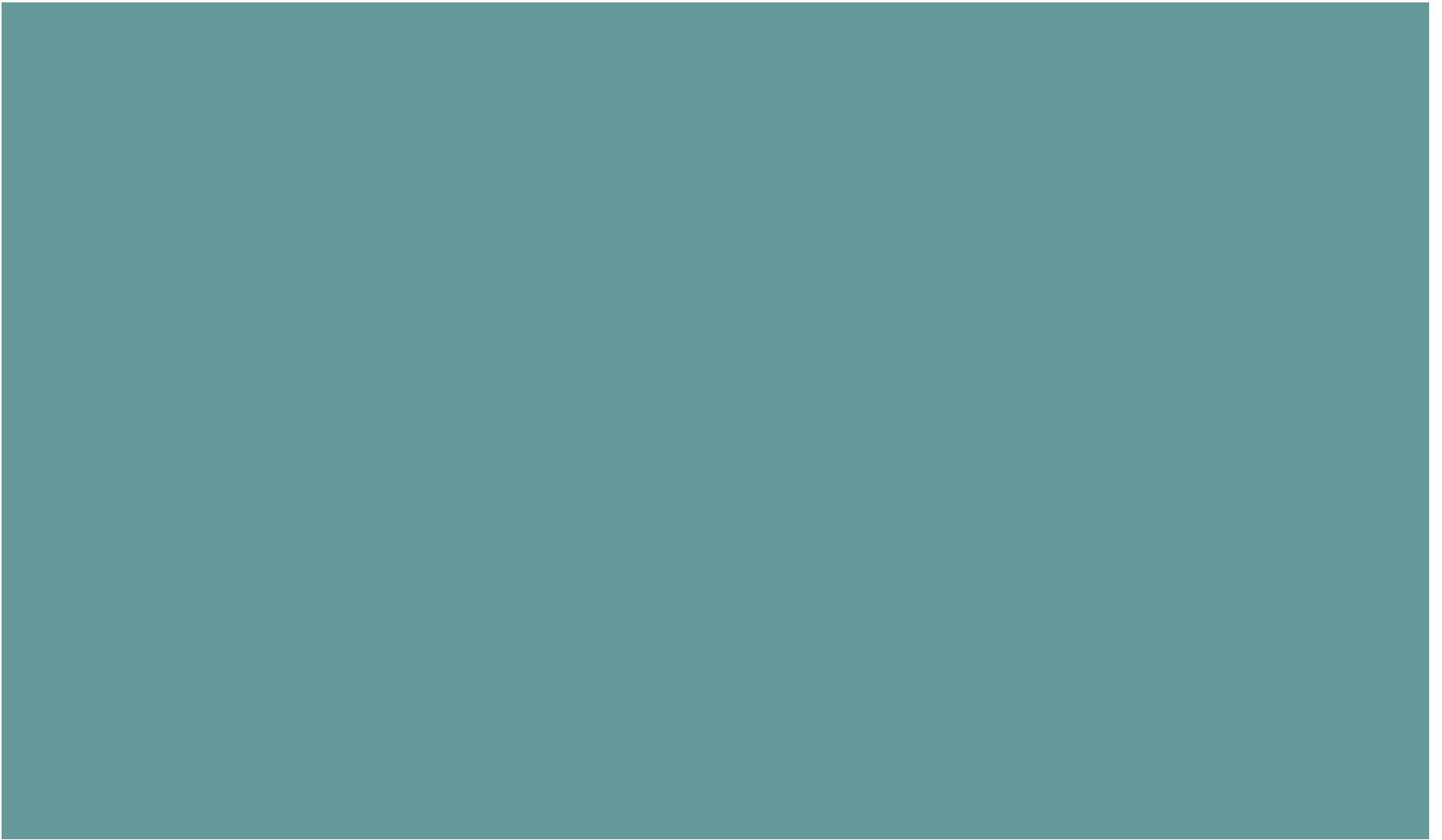
Michigan Basin



Faults

Faults

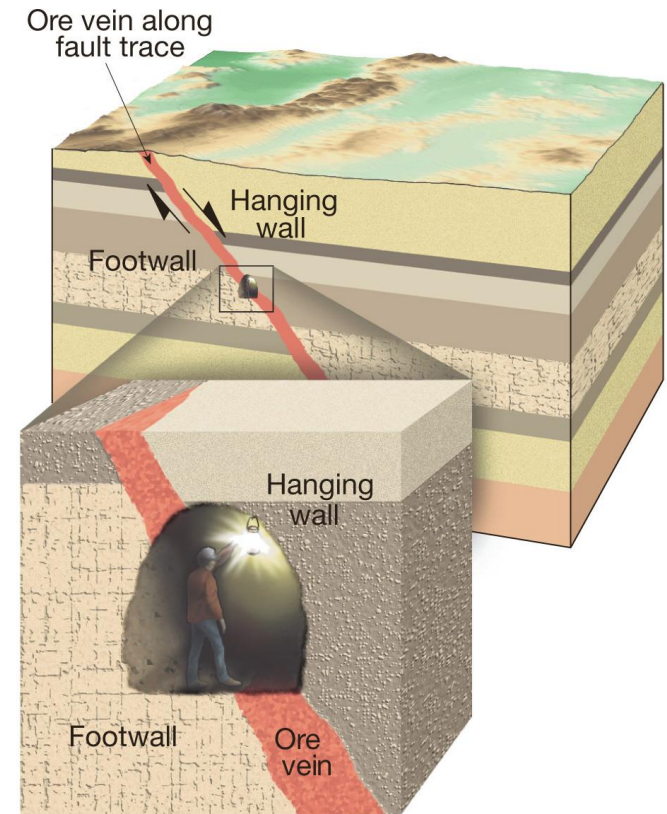
- Faults are fractures with appreciable displacement
- Three types of fault
 - Dip-slip
 - normal fault
 - reverse fault
 - Strike-slip fault



Dip-slip faults

(either a normal or a reverse fault)

- Dip-slip faults have movement parallel to the dip of the fault surface, so they have vertical displacement.
- hanging wall - rock surface above the fault
- foot wall - rock surface below the fault
- fault scarp - line of the fault on the surface, which is generally a long low-lying cliff



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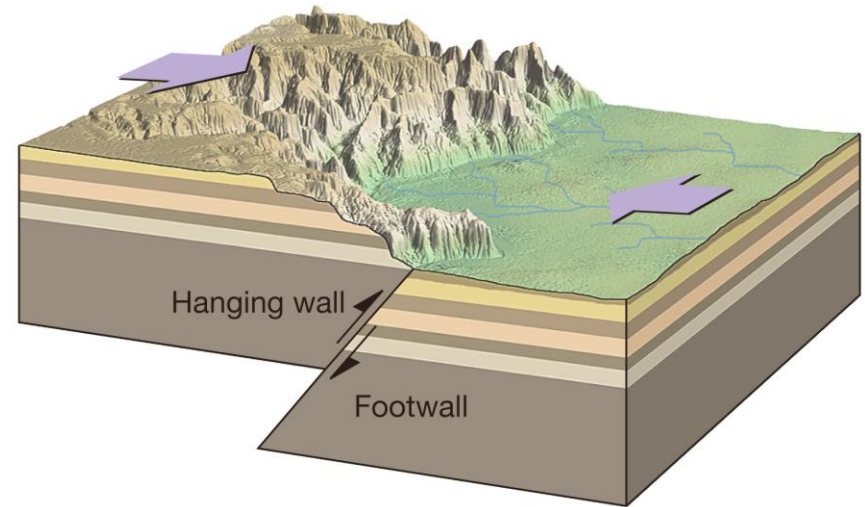
1964 Alaska earthquake



Reverse Fault

(A type of dip-slip fault)

- Hanging wall moves up relative to the footwall
 - older strata ends up over younger strata
- produced by compressional forces (pushing together)
 - results in crustal shortening
- can occur at subduction zones (convergent boundaries)



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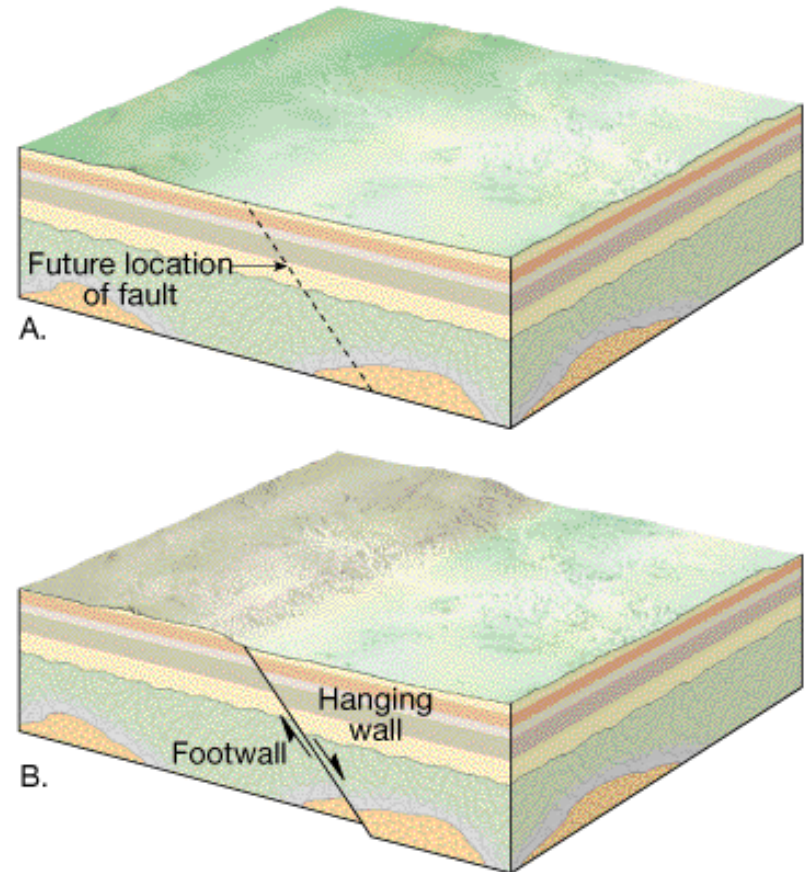
Ketobe Knob, Utah

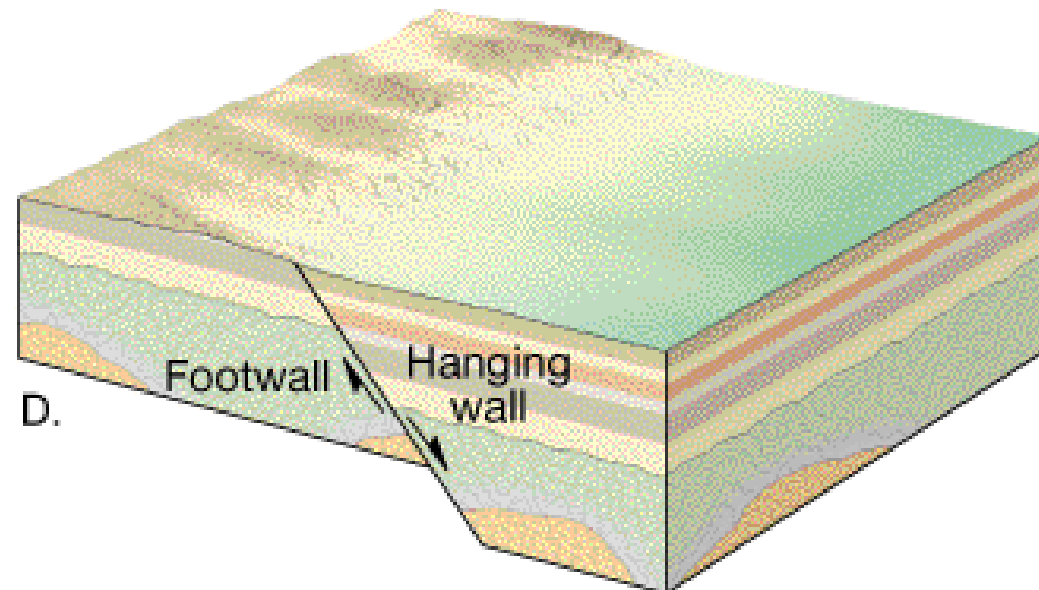
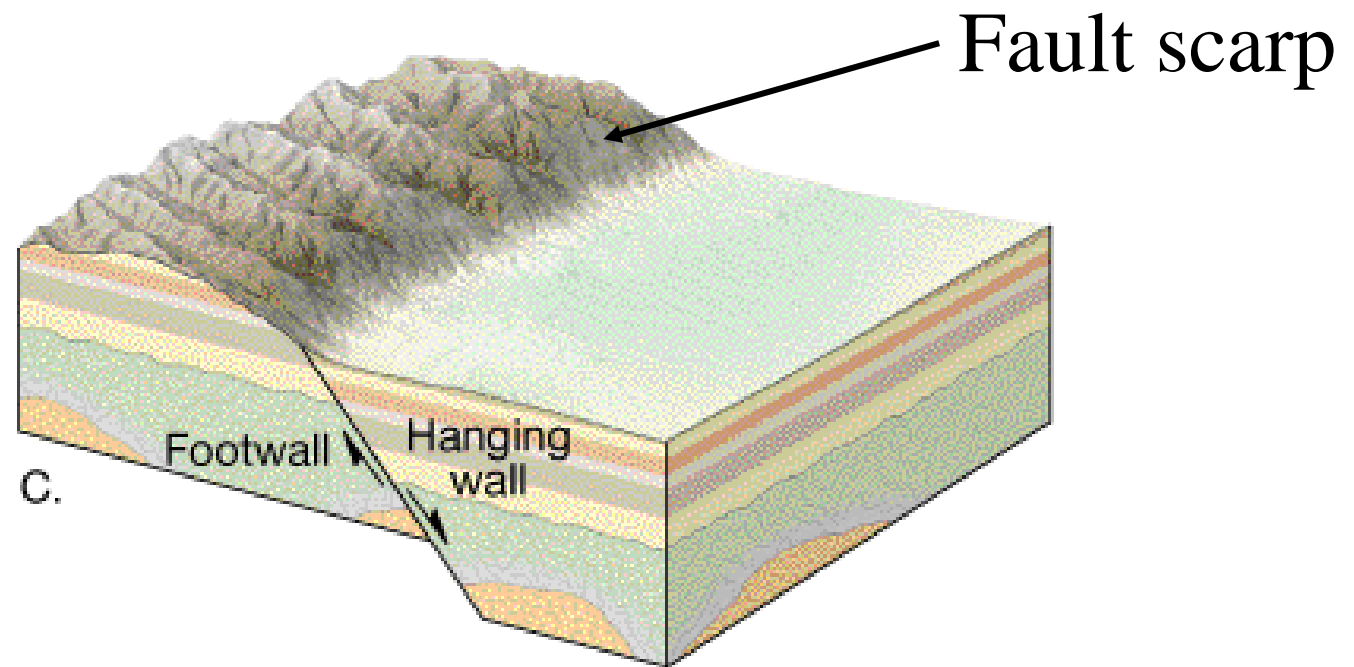
<https://structuralgeo.files.wordpress.com/2013/07/reverse-fault-zone-ketobe-knob-utah.jpeg>

Normal fault

(a type of dip-slip fault)

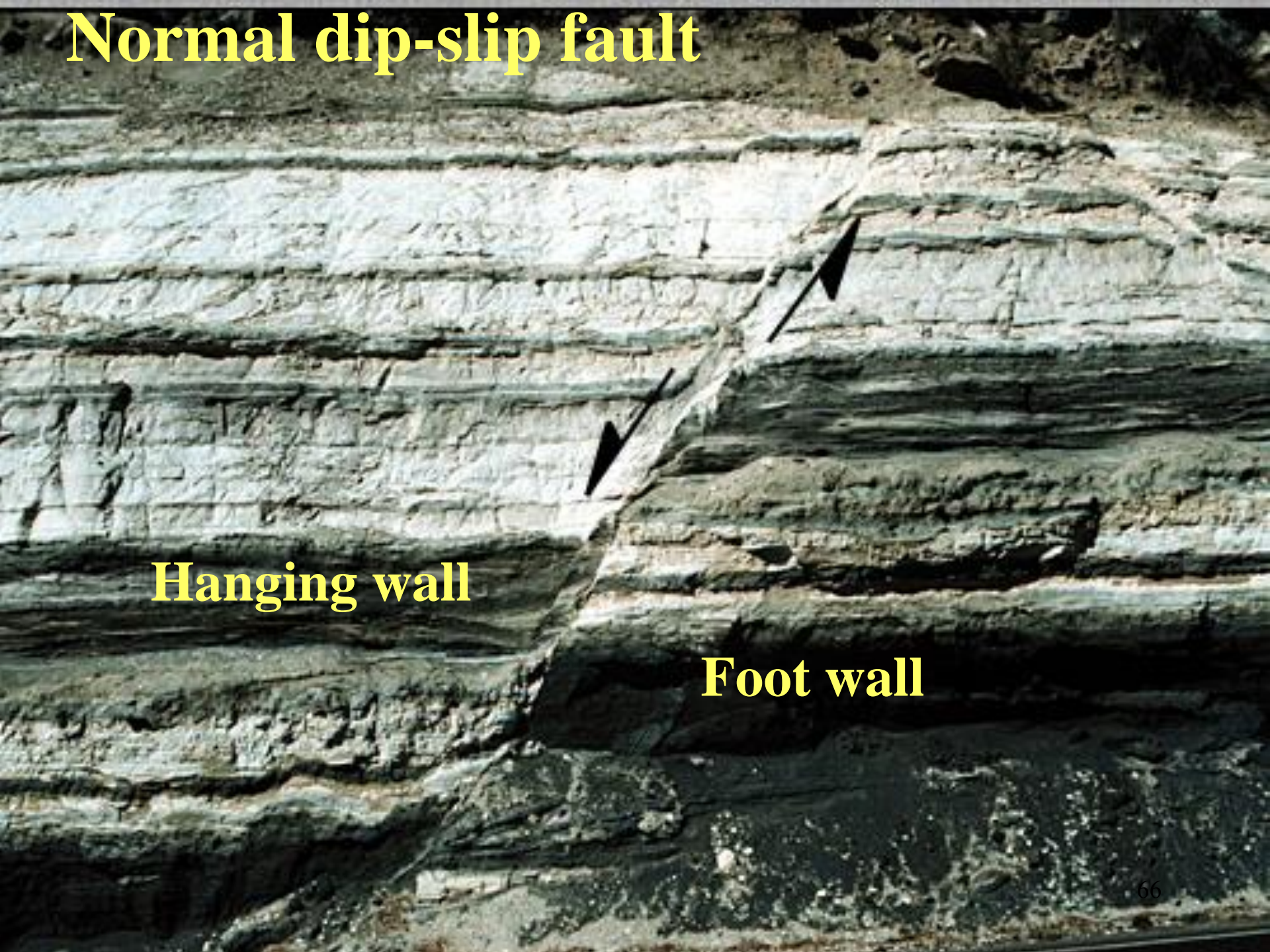
- In normal faults the hanging wall moves down relative to the footwall.
- produced by tensional forces (pulling apart)
 - results in lengthening and extension of the crust
- can occur at spreading centers





Normal fault

Normal dip-slip fault

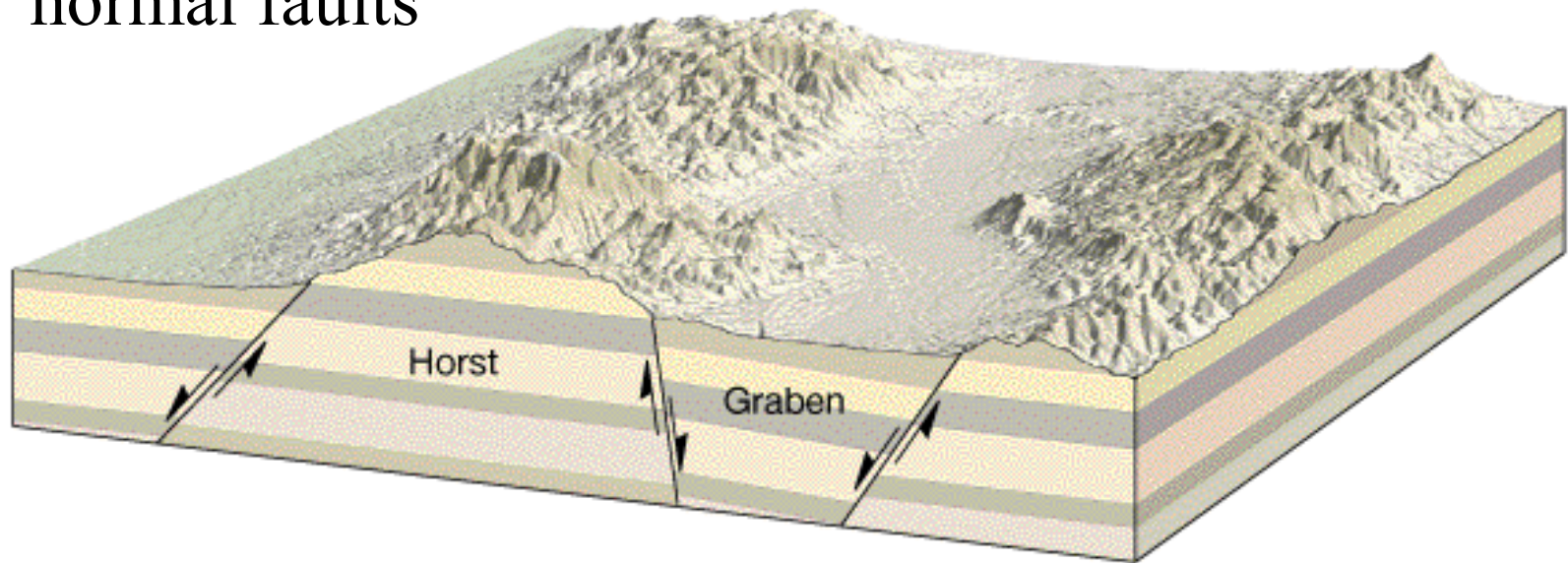


Hanging wall

Foot wall

Features of normal faulted terrain

- horst – uplifted block bounded by two normal faults
- graben – down-dropped block bounded by two normal faults

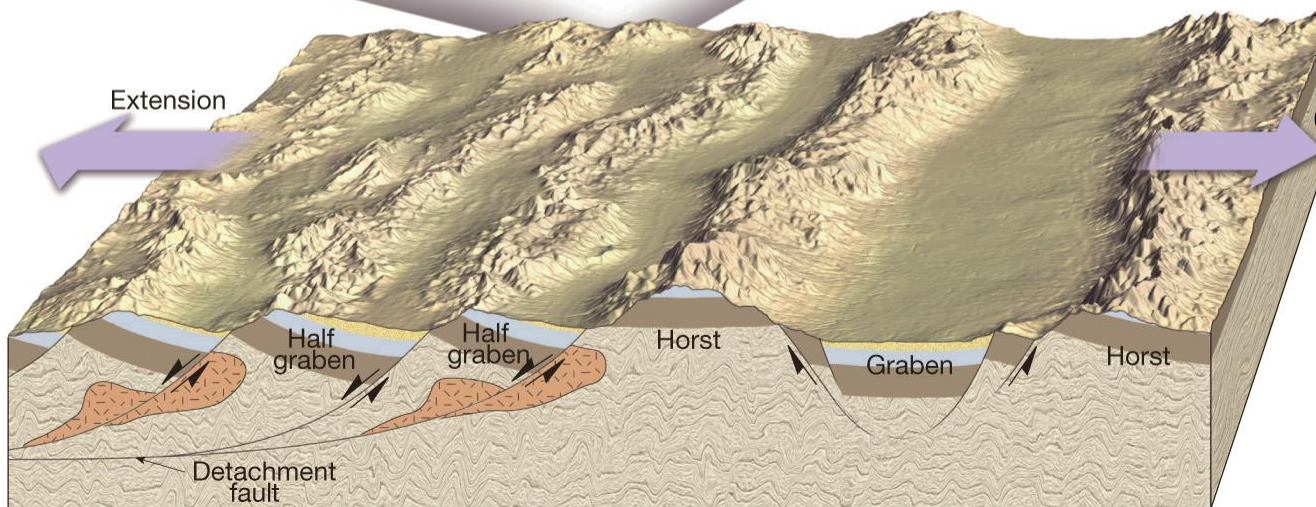


- fault-block mountains are produced in normally faulted terrain.
 - Example - Basin and Range Province of the southwestern United States

Normal faulted terrain: Fault block mountains of the Basin and Range Province of SW United States

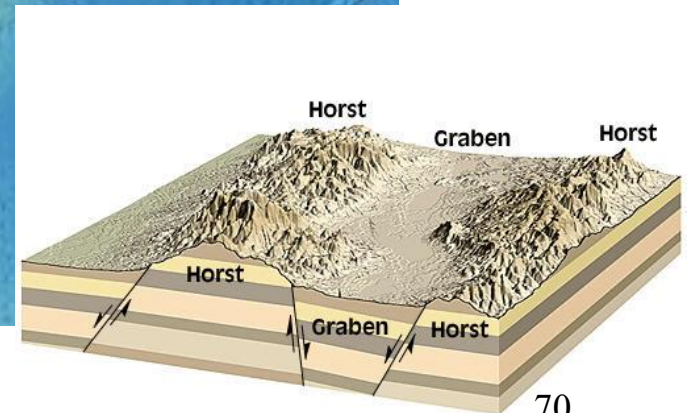
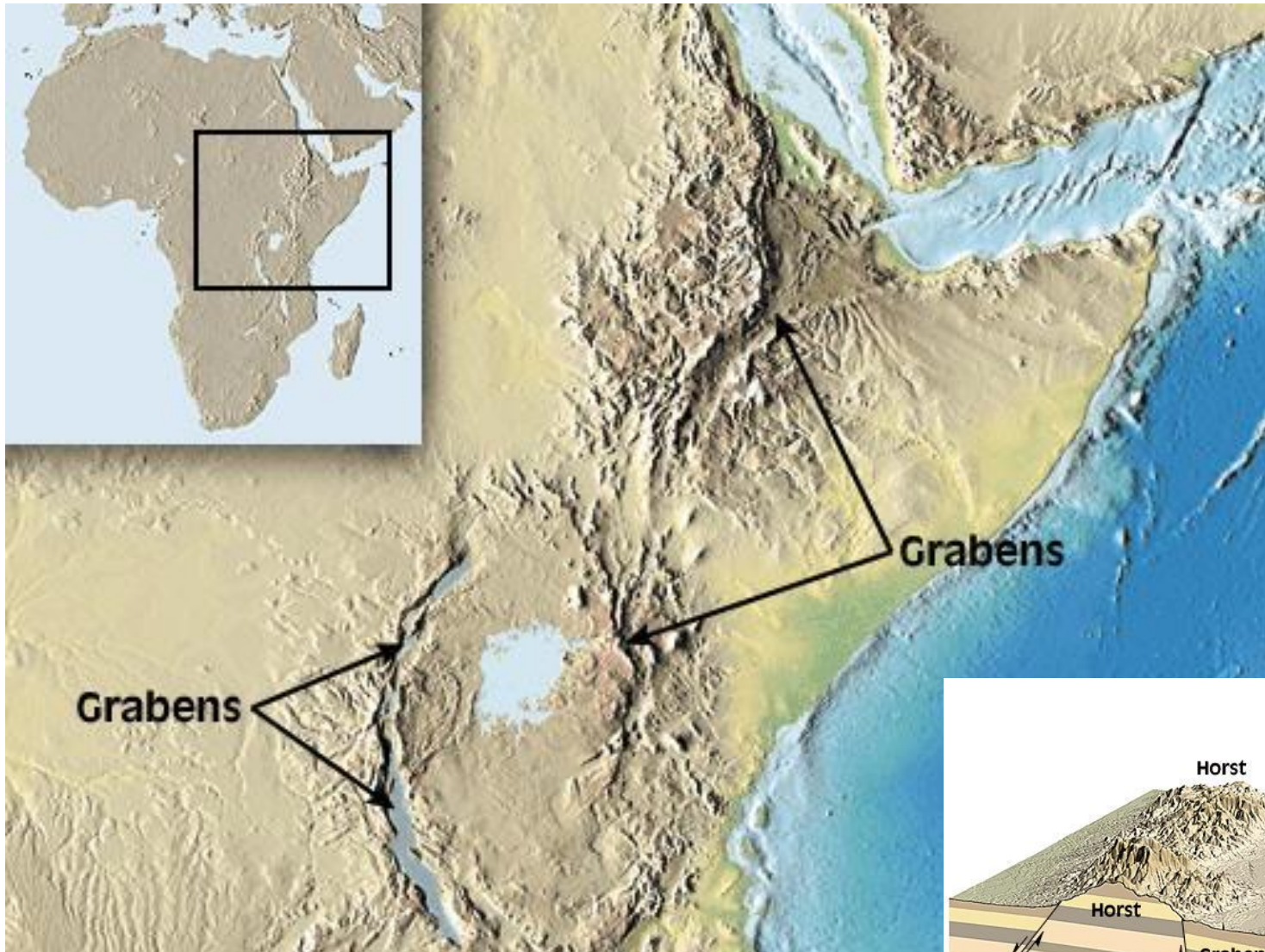


Extensional
forces
lengthen the
crust





Normal faulting occurs at spreading centers



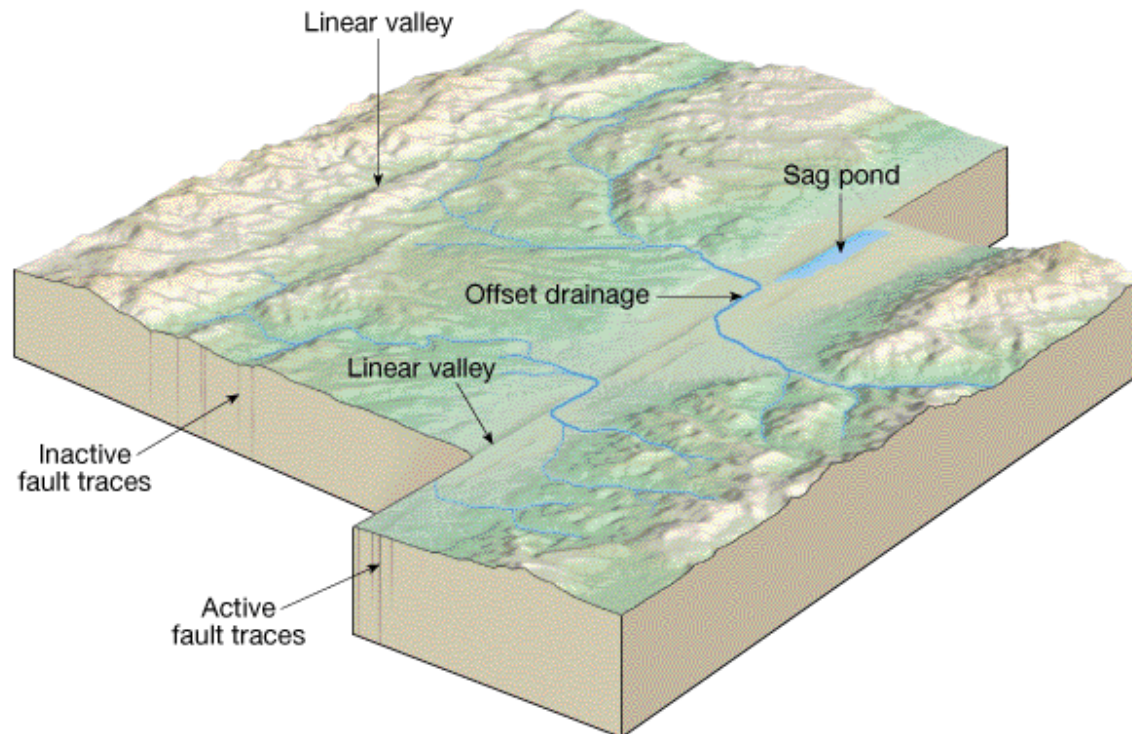
- Clay model of the formation of normal faults at a spreading center. [Movie](#)
- (Clay demonstration by Ron Harris)
<http://geologyindy.byu.edu/faculty/rah/slides/Rock%20Canyon/Wasatch%20Fault/Wasatch%20Fault%20Home.htm>

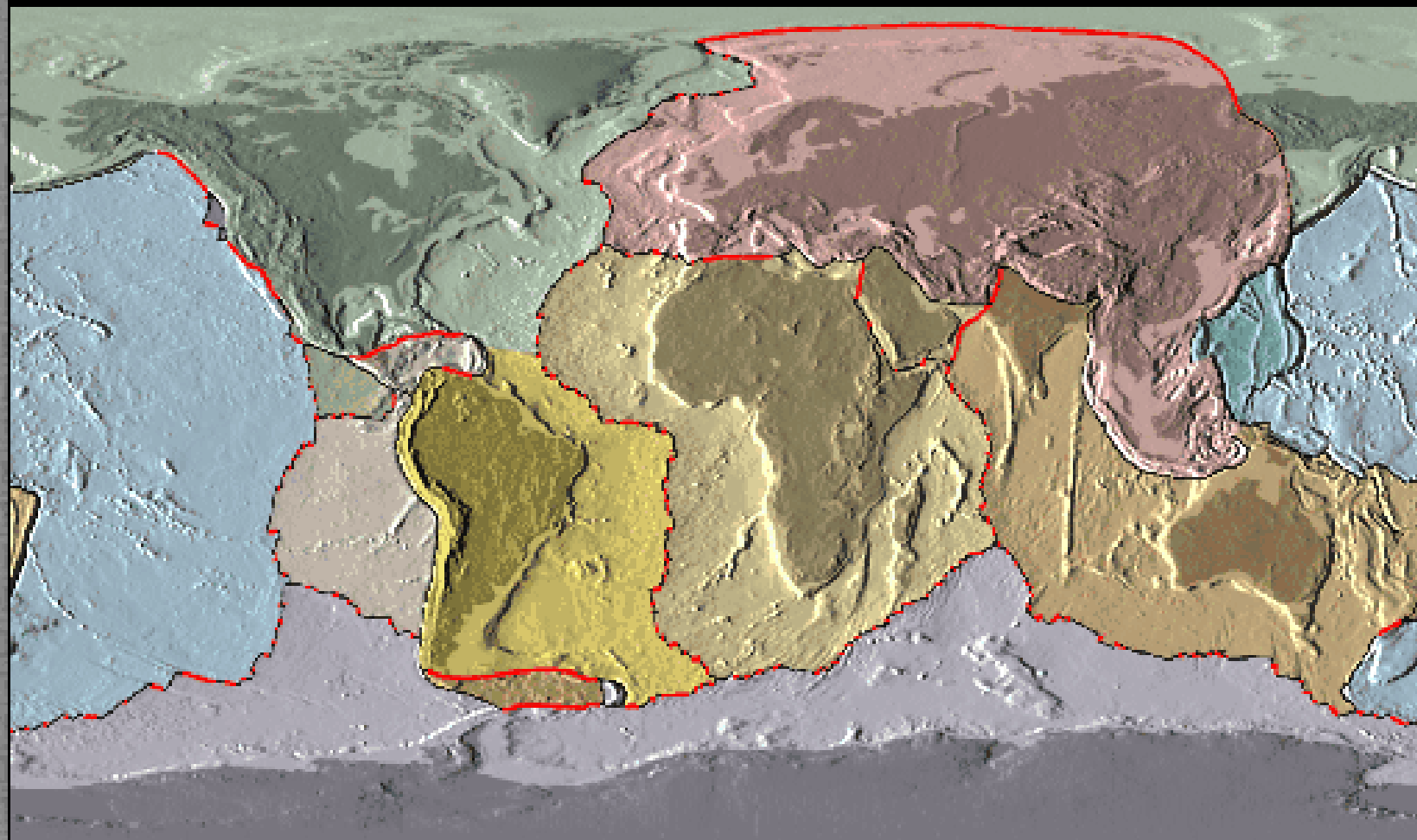
Strike-slip fault

(not a dip-slip fault, no vertical displacement)

Strike-slip fault

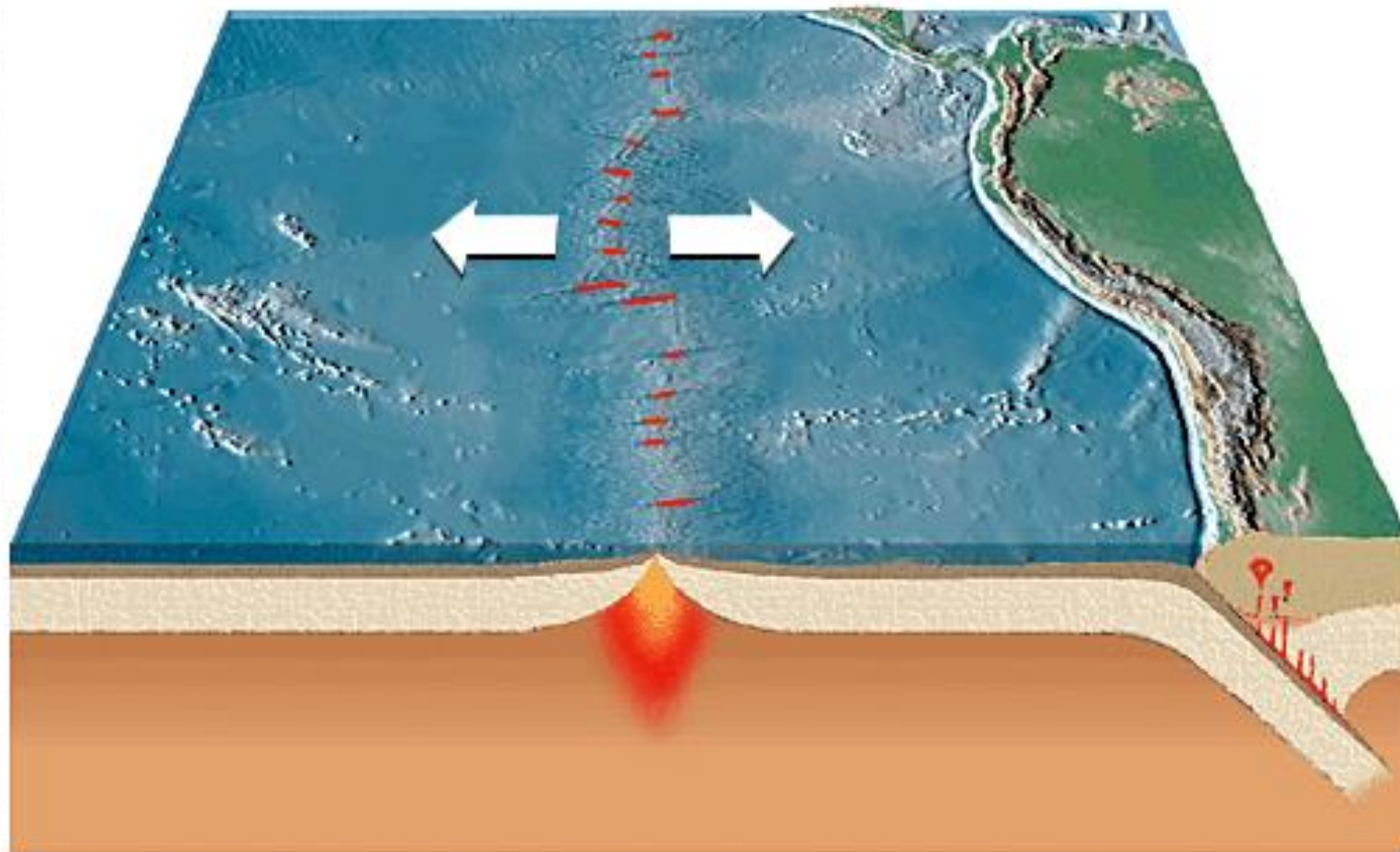
- Fault movement is horizontal and parallel to strike
 - No vertical displacement,
 - fault gouge erodes to produce linear trough
- transform fault are strike slip faults
 - Example: San Francisco fault system (1906 earthquake)





Large strike-slip faults, called **transform faults**, accommodate displacement between plate boundaries.

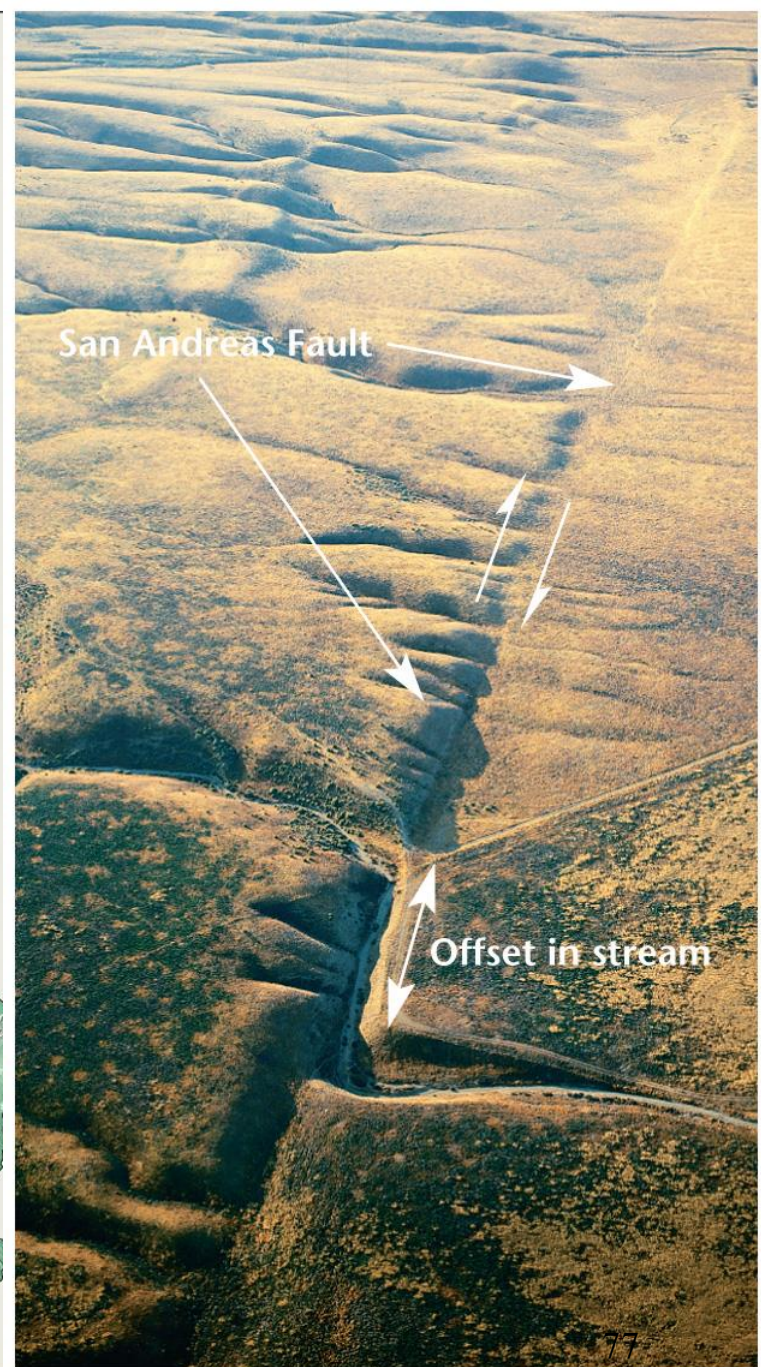
Oceanic ridges

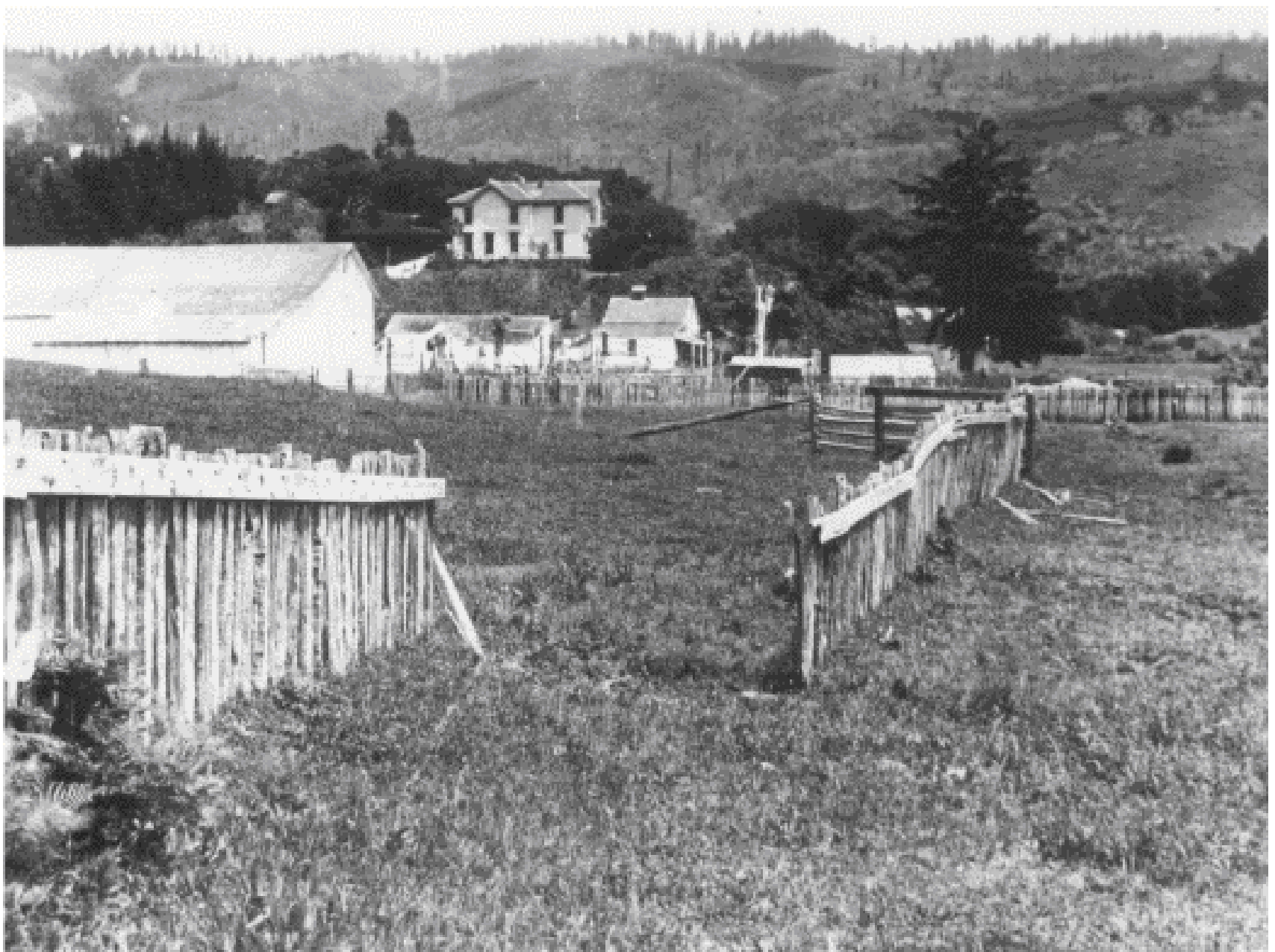


Most transform faults cut the oceanic lithosphere and link spreading centers.



San Andreas Fault System, transform faults

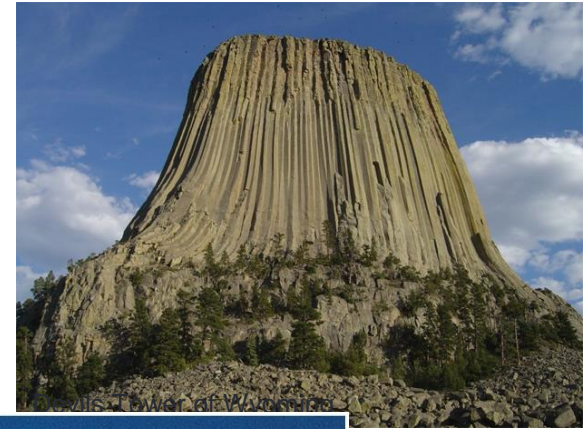




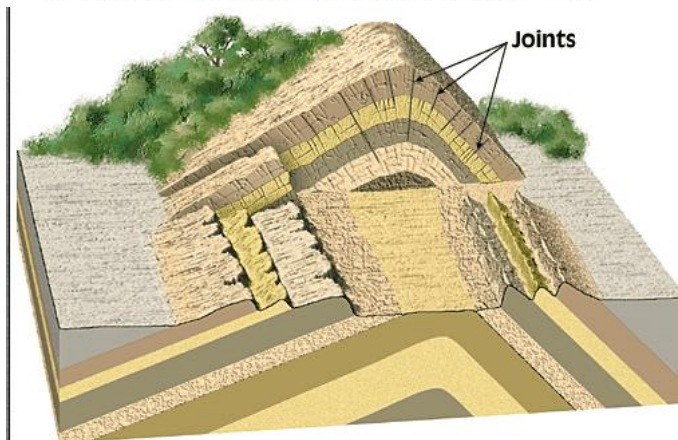
Joints

Joints- are fractures with no appreciable displacement

- Examples of joints
- columnar joints -
 - shrinkage fractures in cooling igneous rock creates elongate pillar-like structures
- sheeting - from unloading of overlying rock
 - joints parallel to the ground surface
 - breaks rocks into “sheets”
 - Creates exfoliation domes
- brittle failure associated with folding



011/08/31/columns-form-



Importance of joints

- Enhanced chemical weathering
- can influence direction of streams
- hydrothermal solutions deposit ore minerals
- risk to construction projects
 - Teton Dam, Idaho, failure 1976
 - seepage along joints

Chemical weathering is enhanced along joints.





End

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